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THESIS

CALCULATION OF ATMOSPHERIC TRANSMITTANCE
BY IBM 3033 COMPUTER CODE LOWTRAN IIIB

by

Moon-Sik Shin

June 1983

Thesis Advisor:

Alfred W. Cooper

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LOWTRAN IIIIB is presently available as a method of predicting atmospheric transmittance at low resolutions at NPS and is suitable for incorporation in simulations and studies of electro-optic weapon/sensor systems performance.



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Calculation of Atmospheric Transmittance
by
IBM 3033 Computer Code LOWTRAN IIB

by

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Submitted in partial fulfillment of the
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ABSTRACT

LOWTRAN IIIB is a FORTRAN computer program for prediction of atmospheric optical transmittance, developed at the U.S. Air Force Geophysics Laboratory (AFGL). LOWTRAN IIIB was received in the modified form developed by Naval Weapon Center China Lake for use on the UNIVAC 1110 computer [Ref. 1], and has now been interfaced to the IBM 3033 computer.

Due to compiler storage limitation in the IBM computer the atmospheric data are read into common storage at the beginning of the program. The two dimensional block data submodule has been replaced with a linear data array, and a new subroutine (array) written to reformat the data. The basic logic structure is unchanged.

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I. INTRODUCTION

LOWTRAN IIIB is a FORTRAN computer program, developed at the U.S. Air Force Geophysics Laboratory (AFGL), which was received in the modified form developed by Naval Weapon Center. It calculates the transmittance of the earth's atmosphere in the spectral region from 0.25 to 28.57 micrometers (350 to $40,000$ cm^{-1}) at 20 cm^{-1} spectral resolution on a linear wavenumber scale. Six atmospheric models, which are tropical, midlatitude summer and winter, subarctic summer and winter, and the U.S. 1962 standard atmosphere, covering seasonal and latitudinal variations from sea level to 100 km are available, in addition to a capability of allowing the user to input atmospheric parameters of his own choosing. The program includes four aerosol models which are average continental, urban, rural, and maritime, and either hazy (5 km visibility) or clear (23 km visibility) haze conditions can be selected in addition to the capability of selecting a particular visibility of the user's choosing. The model accounts for molecular absorption, molecular scattering, and aerosol extinction, plus atmospheric refraction and earth curvature effects.

LOWTRAN IIIB is presently an available method for predicting atmospheric transmittance at low resolution at NPS and may be incorporated in simulations and studies of electro-optic weapon/sensor systems performance.

Basically, this thesis describes a program to

- a) Develop a computer code to calculate atmospheric transmittance for particular wavelengths.

- b) Develop a code to give transmittance and molecular absorptance with bandwidth appropriate to grey-body sources.

c) Apply these to prediction of transmittance over a marine optical path, based on the available meteorological data base for the area, for conditions appropriate to the optical propagation experimental measurement program at the Naval Postgraduate School.

II. THEORY OF ATMOSPHERIC TRANSMITTANCE

Beer's law for linear propagation of monochromatic radiation in the linear regime can be expressed as follows;

$$\frac{d}{dz} [I(\nu)] = -\mu(\nu) I(\nu) \quad (\text{eqn 2.1})$$

or

$$I(z) = I(0) e^{-\mu z} \quad (\text{eqn 2.2})$$

The ratio $\frac{I(z)}{I(0)}$ is defined as the transmittance T of the path of length z , and μ is the total extinction coefficient which is the sum of the coefficients for total absorption and non-forward scattering

$$\mu = \mu_a + \mu_s$$

Both the scattering and absorption coefficients can be divided into components due to the molecules of the air and the aerosol particles suspended in it; i.e.

$$\mu_a = k_m + k_a$$

$$\mu_s = \sigma_m + \sigma_a$$

where k_m = molecular absorption coefficient

k_a = aerosol absorption coefficient

σ_m = molecular scattering coefficient

σ_a = aerosol scattering coefficient

The relative values of the four coefficients depend strongly on the density and molecular composition of the atmosphere and the composition, number density and size distribution of the aerosols. The order of importance in each of the important atmospheric transmission windows is shown in Table I [Ref. 2]. We note from this that scattering by both

molecules and aerosols is of greater importance in the visible, and absorption in the infrared (particularly 8 to 14 μm). Since devices are designed to operate throughout these windows, it is important to predict the transmittance of the atmosphere as a function of wavelength and weather conditions. This prediction becomes a complex problem of computer modelling. The problem requires a definition of the composition, density and pressure of the atmospheric gases, together with the frequencies, line strengths and line widths of all the spectroscopic transitions of the gas molecules and the aerosol constituents, and the number, size

TABLE I
Optical Atmospheric Attenuation Coefficient

Atmospheric window	Wavelength for EO systems (μm)	Attenuation Coefficients in order of importance
Visible	0.4 - 0.7	σ_a , σ_m , K_a
Near Infrared	0.7 - 1.2	σ_a , K_a , σ_m , K_m
Middle Infrared	3.0 - 5.0	K_m , σ_a , K_a , σ_m
Far Infrared	8.0 - 12.0	K_m , K_a , σ_a

and composition distributions of the particles.

Over the entire wavelength range from visible to infrared the absorption by molecules, scattering by aerosols and absorption by aerosols are the dominant extinction mechanisms, and should be considered. In the "window" regions of good transmittance, the molecular line absorption is relatively small, and may in some regions be ignored. The remaining extinction in these regions is due to aerosol scattering (which varies only slowly with wavelength) and "continuum absorption" by the molecules.

A. MOLECULAR ABSORPTION

To compute the monochromatic transmittance of the atmosphere, we must first obtain accurate data describing the frequencies, intensity and line shape of all absorption lines affecting the attenuation. These should be developed a compilation based on certain constraints dictated by atmospheric abundances if we need to develop this monochromatic capability.

The approach by AFGL is to calculate the transmittance at given wavelength for each transition having finite absorption at that wavelength for each of the molecules. The summation over all the molecules gives a monochromatic transmittance which is appropriate directly for laser propagation, but must be degraded by integration over finite bandwidth for low resolution predictions suitable to non-laser sources.

To calculate the transmittance $T = e^{-K_m z}$, the absorption coefficient K_m should be known as a function of frequency for each line.

The four essential line parameters for each line are the resonant frequency, ν_0 (cm⁻¹), the intensity per absorbing molecule, S (cm⁻¹/molecule cm⁻²), the Lorentz line width parameter, \mathcal{L}_0 (cm⁻¹/atm), the energy of the lower state, E'' (cm⁻¹), and the line half-width at half maximum, \mathcal{L} , which is proportional to the pressure. The frequency, ν_0 , is independent of both temperature and pressure. The molecular absorption coefficient is given by

$$K_m(\nu) = \frac{S \mathcal{L}}{\pi [(\nu - \nu_0)^2 + \mathcal{L}^2]} \quad (\text{eqn 2.3})$$

$$S = \int K_m(\nu) d\nu \quad (\text{eqn 2.4})$$

The pressure broadened line width depends in a complicated fashion on temperature; for computation this is approximated by the assumption of temperature - independent collision diameters, leading to

$$\Delta = \Delta(P_0, T_0) \frac{P}{P_0} \left(\frac{T_0}{T}\right)^{\frac{1}{2}} \quad (\text{eqn 2.5})$$

where $P = 1$ atm, and

$T = 296$ degrees Kelvin.

The intensity, S , is pressure-independent and its temperature dependence can be calculated from E'' and

$$S(T) = S(T_0) \frac{Q_v(T_0)}{Q_v(T)} \times \frac{Q_r(T_0)}{Q_r(T)} \times \exp \left[1.439 E'' \left(\frac{T - T_0}{TT_0} \right) \right] \quad (\text{eqn 2.6})$$

where Q_v and Q_r are vibrational and rotational partition functions respectively. At lower pressure the collision broadening diminishes and Doppler broadening becomes important. In the intermediate pressure range the Voigt profile obtained by convolving the two profiles is used. Once Doppler broadening dominates the Gaussian profile may be used. Lorentz linewidths are typically .001 to .01 cm^{-1} , while the Doppler width may be 0.0003 cm^{-1} .

The AFGL calculations [Ref. 14] are based on the Absorption Line Parameters Compilation, which originally listed ν_0 , $S(T)$, $\Delta(P_0, T_0)$ and E'' for 130,000 lines in the range beyond 1 μm , for each one of the species water, carbon dioxide, ozone, nitrous oxide, carbon monoxide, methane and oxygen. A new version of the compilation including over 139,000 lines between 0.2 μm and 30 μm was recently reported by AFGL [Ref. 3].

B. CONTINUUM ABSORPTION

Continuum absorption occurs as a result of collisional interactions between molecules; that is, collisions between two H_2O molecules and those of other gases (principally H_2O : N_2 collision, since nitrogen comprises approximately 80 % of the air)

In wavelength regions between the absorption bands some attenuation occurs of a continuous nature which is attributed to water vapor. The mechanism of water vapor continuum extinction lacks a complete theoretical explanation. At present, it is believed that it results from the accumulated attenuations of the distant wings of H_2O absorption lines, occurring principally in the far infrared part of the spectrum. Other postulates, such as that the phenomenon is caused by other absorption mechanisms involving H_2O dimers, remain possibilities yet to be proved.

However, all that we can do at present is to account for the water vapor continuum phenomenon empirically, based on what limited experimental measurements we have to go on, until better line shape theories become available. It should be emphasized that further accurate and well controlled measurements are urgently required in order to account for this phenomenon in real atmospheric situations with confidence.

A common formulation used (for example in LOWTRAN IIIB) to account for the water vapor continuum attenuation at a fixed temperature has been to define the transmittance $\bar{T}(\nu)$ as follows;

$$\bar{T}(\nu) = e^{-K(\nu) \times \text{range}} \quad (\text{eqn 2.7})$$

where the attenuation coefficient $k(\nu)$ is given by

$$K(\nu) = C_s \left[P_{H_2O} + \frac{C_N}{C_s} (P_T - P_{H_2O}) \right] \omega \quad (\text{eqn 2.8})$$

where P_{H_2O} and P_T refer to the water vapor partial pressure and ambient pressure respectively (atm), and ω defines the quantity of water vapor per unit pathlength (gm cm⁻² km⁻¹). The quantities C_s and C_N are generally referred to as the self and foreign (nitrogen) broadening coefficients for water vapor.

Values for C_s and C_N/C_s have been obtained empirically from laboratory measurements. In LOWTRAN versions I through III, the quantity C_N/C_s is assumed to remain constant over a given wavelength interval. However, a major addition in LOWTRAN IIIB has been to account for the temperature dependence of C_s and this will be discussed in the 8-14 μ m H₂O continuum region. The H₂O continuum radiation in the 3.5-4.2 μ m region is of much less importance and will not be further discussed here.

1. Temperature Dependence

The water vapor continuum attenuation coefficient has been found to have a significant temperature dependence, which was not accounted for in the previous LOWTRAN computer codes. Based on the laboratory measurements using samples of water vapor at elevated temperatures, an approximate empirical expression was obtained by Roberts et al [Ref. 6] for the temperature dependence which is given in Eqn 2.9 below. It was found that the attenuation coefficient due to the water vapor continuum increases as the temperature decreases. That is, for a fixed amount of water vapor in a given path, one would expect more absorption at lower temperatures and less absorption at higher temperatures. This is a somewhat unusual phenomenon. In practice one finds less water vapor in the atmosphere under cold conditions, therefore, the effect of temperature on the attenuation in

the 8-14 μm region plays two competing roles, through the total water content of the path and the attenuation coefficient.

The empirical fits to the wavelength and temperature dependence of the water vapor continuum described in Roberts et al [Ref. 6] have been used in LOWTRAN IIIB with the appropriate conversion of units as follows.

The attenuation coefficient C_s in $\text{gm}^{-1} \text{cm}^2 \text{atm}^{-1}$ at 296 K is given by the following expression in the 8-14 μm region:

$$C_s(\nu, 296) = 4.18 + 5578 \exp(-7.87 \times 10^{-3} \nu) \quad (\text{eqn 2.9})$$

where ν is the wavenumber in cm^{-1} (note that $\nu = 10^4 / \lambda$, where λ is the wavelength in μm).

The temperature dependence of the coefficient C_s was found to vary as:

$$C_s(\nu, T) = C_s(\nu, 296) \exp \left[6.08 \left(\frac{296}{T} - 1 \right) \right] \quad (\text{eqn 2.10})$$

where T is the temperature in degrees Kelvin.

2. Nitrogen Broadened Coefficient

C_N/C_s in the above Equation represents the ratio of the foreign (nitrogen) broadening coefficient to the self broadening coefficient.

In LOWTRAN IIIB we use a value of 0.002 for the parameter C_N/C_s based on the measurements presented by Supplement LOWTRAN IIIB [Ref. 7].

Here, it is assumed that C_N/C_s (at 296 K) does not vary with temperature (since no supporting measurements are available).

Thus, further measurements are needed to determine more accurately the magnitude of the parameter C_N/C_S and its temperature and wavelength dependence.

3. Transmission Calculations

The transmittance due to the water vapor continuum in the 8-14 μm region is calculated for a horizontal path of length RANGE(km) at altitude z using the following expression in LOWTRAN IIIB:

$$T(\nu) = \exp[-C(\nu, 296) \times W(z) \times \text{RANGE}] \quad (\text{eqn 2.11})$$

where $W(z)$ is the effective H_2O absorber amount per unit path length (in $\text{gm} \cdot \text{cm}^{-2} \cdot \text{atm} \cdot \text{km}^{-1}$) at altitude z, and $C_S(\nu, 296)$ is the water vapor (self broadened) attenuation coefficient obtained from laboratory measurements at a temperature of 296 K.

The quantity $W(z)$ is given by:

$$W(z) = w(z) \left[P_{\text{H}_2\text{O}} \exp \left[6.08 \left(\frac{296}{T(z)} - 1 \right) + 0.002 (P_T - P_{\text{H}_2\text{O}}) \right] \right] \quad (\text{eqn 2.12})$$

where

$w(z)$ = $\text{gm cm}^{-2}/\text{km}$ of H_2O in the path at temperature T,
 $P_{\text{H}_2\text{O}}$ = H_2O partial pressure (atm) at altitude z,
 P_T = ambient (total) pressure (atm) at altitude z, and
 $T(z)$ = ambient temperature at altitude z (degrees Kelvin).

Note that the temperature dependence of the attenuation coefficient $C_S(\nu, T)$ given in Eqn 2.10, has been incorporated into the expression for W in Eqn 2.12. The reason for this

is so that the temperature variation over a given atmospheric slant path is weighted equally with the water content along the path.

It may be worth contrasting Eqn 2.12 with the corresponding expression which has been used in LOWTRAN I through LCWTRAN III, that is:

$$W(z) = w(z) \left[P_{H_2O} + 0.005 (P_T - P_{H_2O}) \right] \quad (\text{eqn 2.13})$$

C. AEROSOL EXTINCTION COEFFICIENT & MODELS.

Scattering of radiation occurs from molecules in the air, from aerosol particles suspended in the air, and from water droplets in fog, rain or hail. The attenuation of a beam depends on the size and number density distribution and the refractive index of the particles. Atmospheric transmission may be strongly influenced or even dominated by scattering from aerosol particles. The major parameter determining the interaction is the ratio of particle radius to the wavelength of the radiation.

1. Aerosol extinction coefficient

Aerosol extinction is sum of the absorption and scattering by aerosol particles. Although scattering by aerosol particles whose radius is smaller than about 0.03 times the wavelength of the light may be calculated by Rayleigh theory, for the whole range of particle radius Mie theory must be applied.

a. Rayleigh Scattering

The Rayleigh volume total scattering coefficient may be written as

$$\sigma_m = \frac{8\pi^3(n^2-1)^2}{3N\lambda^2} \left(\frac{6+3P}{6-7P} \right) \quad (\text{eqn 2.14})$$

where n ; refractive index

N ; the number density of molecules

P ; depolarization factor due to some molecular anisotropy.

Following Penndorf [Ref. 8] it can be written in the following form which depends on pressure and temperature.

$$\sigma_m = 4.807 \times 10^{-20} (273/T) (P/1013) \nu^{4.017} \text{ Km}^{-1} \quad (\text{eqn 2.15})$$

The strong dependence on wavelength means that Rayleigh scattering is a small effect for wavelengths longer than the visible.

b. Mie scattering

This is appropriate for the condition of particle size comparable to wavelength; i.e., for large molecules and small droplets. This is the most important mechanism for atmospheric scattering. McCartney gives the scattering coefficient [Ref. 9].

$$\sigma = \pi \int r^2 N_T(r) K(\lambda, n) dr \quad (\text{eqn 2.16})$$

where r ; the particle radius

$N_T(r)$; the total number density of particles in the size range dr about r

$K(\alpha, n)$; the area scattering coefficient, the ratio of effective area to geometrical area, which depends on $\alpha = 2\pi/\lambda$ and refractive index n .

2. Aerosol Model

The range of conditions in the boundary layer is represented by three different aerosol models, which are rural, urban or maritime. The first two are grouped together as "averaged continental".

a. Maritime

Maritime aerosol composition and size distributions are significantly different from rural and urban aerosol types. The maritime aerosol component is due to salt particles which are caused by the evaporation of seaspray droplets. The concentration of particles near the surface is strongly dependent on wind speed (above 7 m/sec) and the size distribution is also dependent on relative humidity. This salt-particle number-density decreases rapidly above about 500 m [Ref. 10].

b. Rural

Rural aerosol background is partly the product of reactions between various gases in the atmosphere and partly due to dust particles picked up from the earth surface. The particle concentration is largely dependent on the history of the air mass carrying the aerosol particles.

c. Urban

Urban aerosols contain certain additives from combustion products and from industry. Shettle and Penn assumed an addition of 35% soot-like particles of similar distribution to the rural aerosol [Ref. 11].

III. LOW RESOLUTION MODELLING

A. MODEL ATMOSPHERES

The LOWTRAN code provides a choice of six atmospheric models. These include the 1962 U. S. Standard Atmosphere plus five supplementary models. Surface level conditions for the supplementary models are given in Table II [Ref. 1]. There are also two conditional haze models - corresponding to sea level visual ranges of 5 and 23 km - provided as basic input data for LOWTRAN III. Aerosol attenuation for other visual ranges is calculated using an interpolation/

TABLE II
Surface Level Condition for Model Atmospheres

Model atmosphere	Latitude & month	Pressure (mb)	Temp. (K)	Air density (g/m ³)	Water vapor (g/m ³)	Ozone (g/m ³)
Subarctic winter	60 N, Jan	1013	257	1372	1.2	4.1E-05
summer	60 N, July	1010	287	1220	9.1	4.9E-05
Midlatitude winter	45 N, Jan	1018	272	1301	3.5	6.0E-05
summer	45 N, July	1013	294	1191	14.0	6.0E-05
Tropical	15 N	1013	300	1167	19.0	5.6E-05

extrapolation procedure which utilizes these two models. In addition to the model atmospheres the user has the option of inserting his own model atmosphere, or of building another model by combining various parts of the six standard models.

Provisions are made in the LOWTRAN program for inserting radiosonde data. There are limits on the accuracy of LOWTRAN transmittance calculations when the input is radiosonde data alone. Radiosondes provide vertical profiles of the synoptic meteorological parameters of temperature,

pressure, humidity, and wind. Information on micrometeorological parameters (i.e., aerosol size distributions, local oxidant concentrations, etc.) is also needed as input. In the absence of micrometeorological inputs LOWTRAN relies on its model atmospheres for that information; this may not necessarily be applicable to a specific location. For further discussion on optical parameters see [Ref. 12].

E. CALCULATION MODELLING

In the application it is impossible to measure transmittance at a single frequency. Instead one measures the transmittance $T(\nu)$ averaged over the spectral bandwidth, $\bar{T}_{\Delta\nu}(\nu)$, accepted by the receiver, as indicated in the equation

$$\bar{T}_{\Delta\nu}(\nu) = \frac{1}{\Delta\nu} \int T(\nu) d\nu \quad (\text{eqn 3.1})$$

where ν is the central frequency in the interval, $\Delta\nu$. Consequently for many applications one is interested in knowing the transmittance of the atmosphere averaged over a relatively wide spectral interval, that is, for low resolution.

Thus the term transmittance is somewhat ambiguous unless it is qualified by some indication of the spectral resolution, $\Delta\nu$, over which it is averaged. This is particularly true in the case of molecular absorption, since the absorption coefficient K_m is a rapidly varying function of frequency. It is because of the rapid variation of K_m with frequency that the averaged transmittance \bar{T} does not, in general, obey the simple exponential law. That is,

$$\bar{T}_{\Delta\nu}(\nu) = \frac{1}{\Delta\nu} \int_{\Delta\nu} \exp[-K_m(\nu) \Delta m] \quad (\text{eqn 3.2})$$

where K_m represents the net monochromatic molecular absorption coefficient. On the other hand the molecular scattering coefficient ($\bar{\sigma}_m$) and the aerosol scattering and absorption coefficients ($\bar{\sigma}_a$ and K_a) are slowly varying functions of frequency, and the average transmittance obeys the simple exponential law provided only the direct transmitted beam is being observed

There are four basic approaches to obtaining a low resolution transmittance value for a given path through the atmosphere due to molecular absorption. These are;

- (1) direct measurements over the required path,
- (2) measurements in the laboratory under simulated conditions,
- (3) line-by-line (monochromatic) calculations based on detailed knowledge of spectroscopic line parameters which are then averaged over the required spectral interval, and
- (4) calculations based on band model techniques (which use available laboratory and/or field transmittance measurements or actual line data as a basis).

From the point of view of computations, method 3 involves a considerable amount of work and computer time, and consequently method 4 has been used most frequently.

Basically, it involves a graphical or computerized "curve fitting" technique for accessing the stored transmittance data for the appropriate absorber and scatterer amounts computed from the input atmospheric path conditions. The predetermined variations of transmittance with frequency are stored for the various atmospheric constituents, for standard path conditions.

For a given set of meteorological conditions and selected path, the appropriate absorber and scatterer amounts in the required path are computed for each component and the results used to correct the transmittances for these components. The transmittances for the separate processes

(line absorption, continuum absorption, molecular scattering, aerosol scattering and aerosol absorption) are then multiplied together to give the overall absorption. That is :

$$T_{\Delta\nu}(\nu) \text{ (total)} = T_{\Delta\nu}(\nu) \text{ (line absorption)} \times \\ T_{\Delta\nu}(\nu) \text{ (continuum absorption)} \times \\ T_{\Delta\nu}(\nu) \text{ (Rayleigh)} \times \\ T_{\Delta\nu}(\nu) \text{ (aerosol)}.$$

It should be noted that the LOWTRAN computer code is designed to calculate the transmittance for spectral bands. It should not be used to calculate transmission for laser lines. Suitable techniques for computing atmospheric transmission for laser lines or extremely narrow spectral bands are described in [Ref. 13], [Ref. 14], and [Ref. 15].

IV. EXAMPLE OF PROGRAM USE

A. PROBLEM

As a check on the performance of the code a sample calculation was carried out using the same input data described by Shlanta and Cornette [Ref. 1], for which they provide computed output.

The example is the calculation of the transmittance from 2350 to 2450 cm^{-1} in steps of 5 cm^{-1} for a slant path from 2.5 km to 8.5 km at a zenith angle of 65 degree, for a subartic winter model atmosphere, and a 23 km visual range.

The program is initiated with data input on 4 cards (lines) defining the conditions of the computation, and user directions are given at the beginning of the program listing (Appendix A).

```
CARD 1  **5**1**1
        5 ; Sub-artic winter atmosphere
        1 ; An average continental aerosol model
        1 ; For 23 km visible range
        0 ; For normal operation, etc.

CARD 2  **2**0*****2.500*****8.500*****65.000
        2 ; Vertical & slant path between two altitudes
        0 ; For normal operation of the program which
           selects the shorter path when applicable
        2.500 ; Observer altitude (km)
        8.500 ; Scurce altitude (km)
        65.000 ; Zenith angle at H1 (deg.).

CARD 3  **2350.000**2450.000*****5.000
        2350.000 ; Initial frequency (cm-1)
        2450.000 ; Final frequency ( " )
        5.000 ; Frequency intervals at which
                transmittance is printed.
```

CARD 4 **0

0 ; To end data.

B. OUTPUT FROM LOWTRAN IIIB.

The output for this problem is given in Appendix C. The parameters defining the atmospheric path, model atmospheres and frequency range are first printed out. Following tabulations give the absorber amounts for horizontal and vertical path. At the heading HORIZONTAL PROFILES there are 13 columns. The first column gives a running integer associated with each level. The second column gives the level altitude in km. The next 8 columns give the equivalent absorber amounts per km for the following absorbing species: water vapor, uniformly mixed gas, ozone, nitrogen continuum, water vapor continuum (10 μ m), molecular scattering, aerosol extinction and UV ozone, respectively. The next three columns give the mean refractive index modulus from that level to the level above, the equivalent absorber amounts per km for the water vapor continuum (4 μ m) and for nitric acid.

A heading VERTICAL PROFILES is then printed followed by 15 columns. The first and second columns give the integer associated with the levels traversed by the path and the height of the level. Then follow 8 columns which give the integrated equivalent absorber amounts from the initial altitude to the level above (in the same order as indicated above). The next 4 columns are labelled PSI, PHI, BETA, and THETA (see Appendix D).

The total equivalent absorber amounts for each absorber species are then summarized below in their appropriate units.

The second line in the total equivalent absorber amount table gives the water vapor continuum amount ($4 \mu\text{m}$) and the nitric acid amount.

A transmittance table, containing 12 columns, now follows. The first 3 columns give the frequency (cm^{-1}) wavelength (μm), and total transmittance. The next 7 columns show the individual transmittances due to water vapor, uniformly mixed gases, ozone, nitrogen ($4 \mu\text{m}$) continuum, total water vapor continuum, molecular scattering, and aerosol extinction. The last 2 columns give absorption due to aerosols and the cumulative integrated absorption. The latter quantity can be used to determine the average transmittance over any given spectral interval within the spectral range covered by the calculation. Finally, the total integrated absorption from $V1$ to $V2$ is printed out together with the average transmittance over the band.

V. APPLICATION

The purpose of this implementation of LOWTRAN IIIB is to present a simple method of predicting atmospheric transmittance (at low resolution) which is applicable over a wide spectral interval and for a wide range of atmospheric path.

In this study the bandwidths of the computations have been chosen to match the transmission bands of the filters used by the atmospheric optics measurement group at NPS [Ref. 16]. These filters have been used with grey-body sources to give wavelength resolution in a number of trans-

TABLE III
Broadband System

center wave- length (μ m)	filter bandwidth at $1/2$ ht (μ m)	" (CM-1)	grey-body source temp. (K)	detector
0.49	.010	20619-21053	2800	Si
0.63	.010	15748-16000	"	"
0.84	.010	11834-11976	"	"
1.03	.010	9662-9756	"	"
1.06	.010	9390-9479	"	"
1.60	.098	6064-6447	"	"
2.15	.097	4549-4759	"	Ge
3.80	.400	2500-2778	1800	InSb, 77K
3.835	.110	2571-2646	"	"
10.66	2.850	827-1083	"	HgCdTe, 77K
11.02	.710	879-938	"	"

mittance measurements at Monterey and elsewhere. Figure 5.1 shows the passbands of the grey-body source filters superimposed on curves of atmospheric transmittance for a 1000 ft, path at sea level containing 5.7 mm of precipitable water at 79°F. The filter bands are shown as Table III.

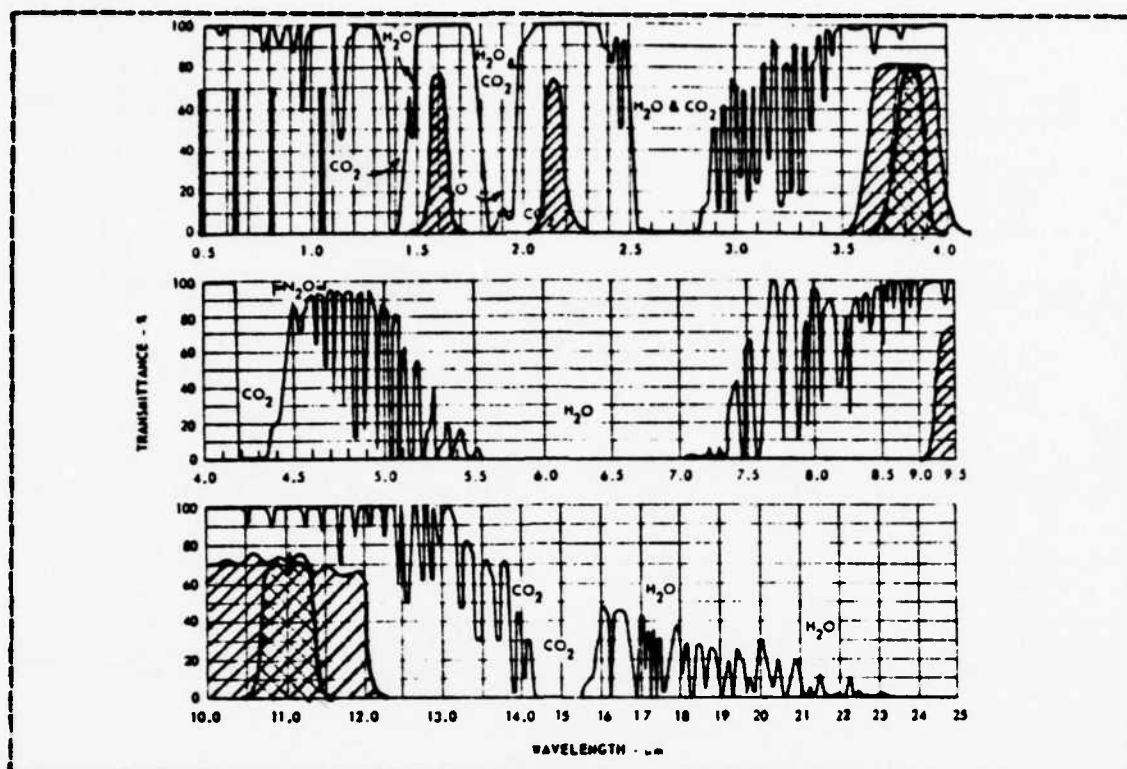


Figure 5.1 Bandpass Regions for Grey-body Source.

A. LOWTRAN CALCULATIONS FOR THE MONTHLY AVERAGE DATA OF MONTEREY BAY

Calculations intended for use in the propagation studies in the Monterey Bay Area make use of the known range of conditions and three different relative humidities, 70, 80, and 90 %, taking into account the monthly and yearly averages of air pressure, air temperature, dewpoint temperature, water vapor density, and relative humidity. These monthly and yearly averages are taken from the thesis by Guner [Ref. 4], and ozone density is translated from that source to $6.00\text{E-}05 \text{ gm/m}^3$, consistent with the midlatitude model in Table II. In this Table IV the average temperature is used which is the mean value of maximum and minimum average temperatures corresponding to day/night conditions.

Table V shows the calculated atmospheric transmittance for 3 different cases of relative humidities and temperatures for selected wavebands of the broadband system at sea level (1.60, 2.15, 3.80, 3.835, 10.66, 11.02 μm), 13.16 km range, and with typical May conditions. Figure 5.2 and Figure 5.3 also represent the atmospheric transmittance of Table V.

By comparing the numbers for different relative humidity, it can be seen that increasing relative humidity

TABLE IV
Monterey Bay Area Average Weather Condition

	mean pr. (mbars)	mean temp. (deg.)	mean dew. (deg.)	H ₂ O den. (gm/M)	rel. hum. (%)
JAN.	1020.0	10.03	6.12	7.37	77
FEB.	1020.0	10.57	6.68	7.67	76
MAR.	1017.5	10.84	6.68	7.67	78
APR.	1018.0	11.68	8.34	8.53	77
MAY.	1017.0	13.34	9.45	9.11	77
JUN.	1015.0	14.45	11.12	10.16	79
JUL.	1017.0	14.73	11.68	10.55	83
AUG.	1017.5	15.29	10.23	10.92	84
SEP.	1013.5	15.84	14.73	10.55	79
OCT.	1017.5	14.73	10.01	9.40	75
NOV.	1019.0	12.79	7.23	7.95	72
DEC.	1018.0	11.12	6.12	7.37	73
YEAR.	1017.5	13.07	8.89	8.82	77

TABLE V
Comparison of Transmittance by R.H & Temperature

	μm	wavenumber cm ⁻¹	Rel. Humidity			temperature (°C)		
			70 %	80 %	90 %	0.0	13.3	30.0
(1)	1.60	6064-6447	.823	.823	.823	.823	.823	.823
(2)	2.15	4549-4759	.834	.832	.831	.833	.833	.833
(3)	3.80	2500-2778	.803	.797	.791	.794	.799	.803
(4)	3.835	2571-2646	.813	.808	.803	.806	.810	.814
(5)	10.66	827-1083	.836	.815	.793	.801	.821	.839
(6)	11.02	879-938	.856	.833	.808	.818	.840	.860

decreases the atmospheric transmittance. We can also see that increasing temperature increases the transmittance;

significant differences of the transmittance occur due to temperature. The numbers from the 8-14 μm region are much larger than the other band. This point is consistent with the theory of the temperature dependence from chapter II.

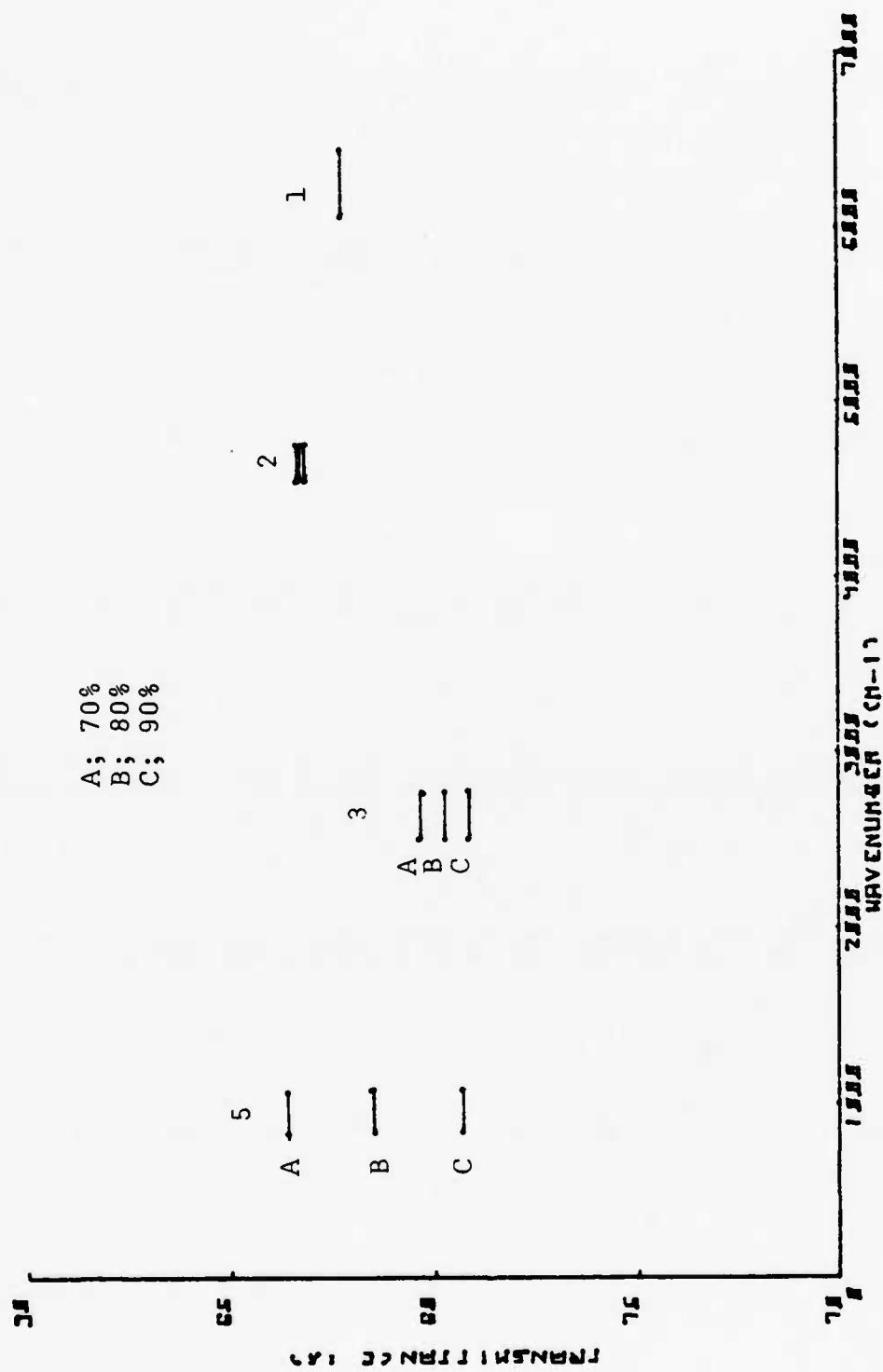


Figure 5.2 Comparison of Transmittance due to Relative Humidity.

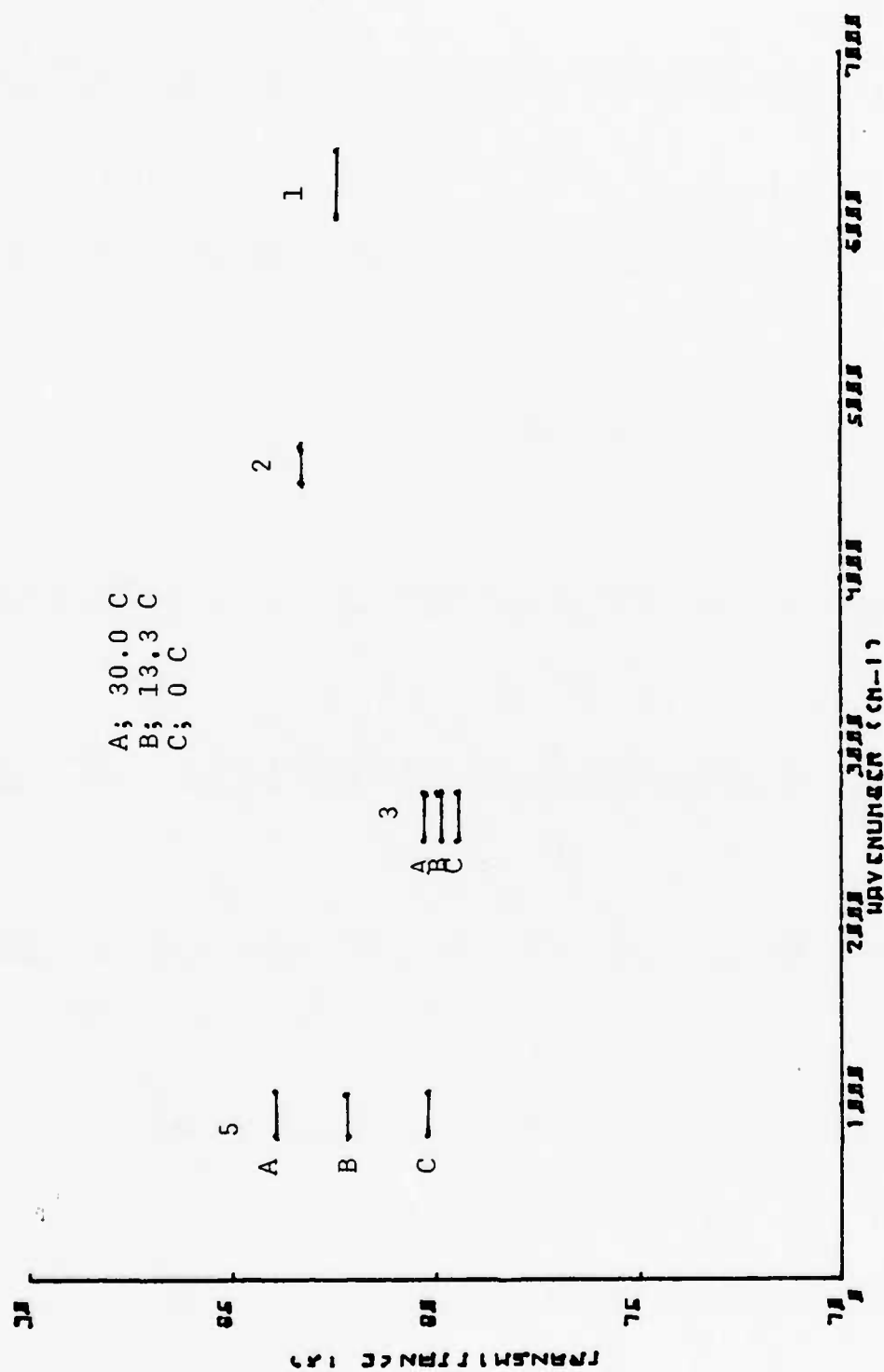


Figure 5.3 Comparison of Transmittance due to Temperature.

B. OPTICAL TRANSMISSION AT SAN NICOLAS ISLAND, CEWCOM-78.

In May 1978 the NPS Optical Physics and Micrometeorology Groups participated in a Cooperative Experiment, West Coast, Oceanography and Meteorology, 1978 (CEWCOM-78) in cooperation with the measurement program OSP-III run by the USN Optical Signatures Program on San Nicolas Island. Optical transmittance measurements from the R/V ACANIA to shore and micrometeorological measurements on the R/V ACANIA were made by NPS measurement group, and other meteorological data were monitored in the OSP experiment. Weather conditions limited the data obtained.

The measured transmittance of the optical path from the R/V ACANIA to site A on the northwest tip of San Nicolas

TABLE VI
San Nicolas Data, CEWCOM-78

Date	Time	Range (m)	Broadband (μ m)	
			3.6-4.0	9.0-12.0
5/15	1126		63.0 % (94.99)	
"	1458	2192		45.2 % (96.14)
"	1633	2200	57.8 % (95.31)	
"	1649	2200		49.8 % (96.63)

Island is listed in Table VI. The transmittances are the observed values, uncorrected for the molecular absorption, measured with filter selection from a grey-body source.

The micrometeorological data taken by the NPS Meteorology Team during the CEWCOM-78 experiment [Ref. 16] and the Daily Statistical Summary of Meteorological Data [Ref. 17] are used to find interpolated temperature, relative humidity, and visibility. The numbers in parentheses represent LOWTRAN predictions of transmittance, including both molecules and aerosol extinction based on these meteorological input data.

The comparison is given for the one day for which suffi-

TABLE VII
Meteorological Data

Time	Broadband (μm)	Press. (mbar)	Temp. (deg.)	Dew T. (deg.)	R.H. (%)	Water vapor (g/m ³)
1126	3.6-4.0	1013.2	14.52	10.0	81.0	9.8
1633	"	"	14.99	10.9	66.7	"
1458	9.0-12.0	"	14.80	10.9	74.67	9.7
1649	"	"	14.93	10.0	65.69	9.5

cient meteorological data are available (Table VII). This comparison shows the LOWTRAN predicted transmittance for these conditions to be much higher than the measurements. The overestimation of transmittance by LOWTRAN suggests that the particle model included in the code may be inadequate, or that the model's total particle number density may be underestimated by the code. Calibration errors in the measurement equipment have been investigated and are considered an unlikely source of the discrepancy.

C. EXPERIMENTAL MEASUREMENT FROM MARINA TO POINT PINOS

The current modelling effort was undertaken in support of an experimental program of transmittance measurement of total extinction on ranges over Monterey Bay from Marina to Point Pinos (June 1979) using laser and broadband (thermal) sources. The optical extinction was measured at 0.4880, 0.6328, 1.03, 1.06, and 11.05 μm .

The calculational program has therefore been directed to computation of molecular absorption and aerosol extinction coefficients for sea level propagation under a range of weather conditions statistically typical of Monterey. This has been done for a variety of wavelengths appropriate to the measurements, and averaged over the filter bandwidths for the broadband sources used (see Table III.).

The aerosol particle spectrum was measured on board the ACANIA with a Knollenberg counter. This information was used to predict a scattering extinction coefficient, based on one point measurement of the particle size spectrum.

For comparison of data it is necessary to subtract the effects of molecular absorption from the total optical

TABLE VIII
Extinction Coefficient from Optical Measurement and LOWTRAN

date time		optical		LOWTRAN		diff.	aeroscl
		(nm)		total	molecular		
		(a)	(b)	(c)	(a-c)	(d)	
6/6	1734	1.60	22.5	2.47	0.072	22.4	31.6
	1714	1.06	21.7	2.68	0.165	21.5	32.4
	1718	1.03	19.4	2.69	2.63	16.8	32.9
	1720	0.84	22.1	2.75	0.196	21.9	32.8
	1721	0.63	22.5	2.83	0.066	22.4	32.8
6/7	1406	11.02	19.17	1.84	1.52	17.7	7.62
	1404	10.6	12.1	2.01	1.49	10.6	7.72
	1423	1.06	35.5	2.45	0.16	35.3	48.7
	1434	1.03	43.9	4.87	2.56	41.3	49.5
	1555	11.02	20.0	1.73	1.25	18.8	7.83
	1558	10.6	10.6	1.77	1.26	9.34	7.71
	1615	1.6	35.3	2.17	0.07	35.2	49.2
	1628	1.06	38.3	2.43	0.15	38.2	50.5
	1632	1.03	35.0	4.69	2.4	32.6	51.4
	1634	0.84	40.1	2.52	0.18	39.9	51.0
	1636	0.63	43.2	2.45	0.067	43.13	53.0

extinction. This is done using LOWTRAN. The value of relative humidity and water vapor content have been obtained from Prof. Schacher and Dr. Fairall who made the aerosol measurements [Ref. 18].

The results are presented in Table VIII. The extinction in units of $E-02 \text{ km}^{-1}$ as measured optically, the LOWTRAN prediction of total extinction, the molecular contribution computed by LOWTRAN, total measured minus LOWTRAN molecular, and the values calculated from aerosol spectra are shown.

By the comparison seen by the Table VIII, optical extinction computed from the aerosol spectra is much higher than the LOWTRAN total extinction, and is very often higher

than the measured optical extinction. This may be due to the fact that aerosol extinction was not measured over the whole path but at only one point. It was measured on board the ACANIA meteorological mast at a height of approximately 10 m above mean sea level, at approximately the centerpoint of the optical path. The optical path in this experiment ranged from 18.3 m above sea level at Marina to almost sea level at the ACANIA, due to earth curvature over the 13.16 km path.

Total aerosol extinction from LOWTRAN is much less than the extinction calculated from the measured aerosol spectrum. Subtracting the calculated molecular component of the LOWTRAN from the optically measured extinction yields the aerosol extinction ("diff."), which is much greater than the LOWTRAN aerosol component but somewhat less than that predicted from the aerosol spectra. i.e., agreement of aerosol extinction is much better with aerosol spectrum measurement than with the LOWTRAN aerosol extinction. It appears probable that the LOWTRAN model consistently underestimates the severity of aerosol extinction in the marine environment.

VI. CONCLUSION AND RECOMMENDATION

The comparisons of LOWTRAN IIIB predictions with optical measurements and with computations from particle size distribution measurements show considerable differences. The differences could conceivably arise from measurement errors in the optical transmittance and the particle measurements; the two physical measurements give results reasonably close to each other but in both cases differ markedly from LOWTRAN IIIB. In view of the consideration given to both measurement techniques it seems most likely that LOWTRAN IIIB consistently underestimates the aerosol attenuation in the infrared wavebands.

It must first be realized that LOWTRAN IIIB contains limitations. It does not include the effects of

1. clouds, fog or precipitation
2. refraction, scintillation or distortion of the propagating beam due to turbulence
3. background interference from direct or scattered solar radiation
4. surface level reflections and masking
5. multiple scattering

In addition the models of refraction and earth curvature are simplified and the atmosphere is considered to be horizontally homogeneous and stable with no inversion. These limitations mean that at best LOWTRAN IIIB results are of "moderate" accuracy.

Some of these limitations have been removed in the continuing evolution of LOWTRAN. LOWTRAN IV (Feb. 1978) in addition to using improved empirical transmittance functions for gases of small absorber amount, includes the effects of nitric acid as an absorber and emitter, and incorporates a

"radiance mode" to compute the radiance from the earth surface and the atmosphere.

LOWTRAN V (Feb, 1980) has been upgraded to include new aerosol models dependent on altitude and relative humidity, and a model for fog attenuation.

It may be expected that implementation of these codes (preferably LOWTRAN V) may give better agreement between experiment and computation and allow better validation of aerosol extinction calculation methods.

LOWTRAN IV and LOWTRAN V are not currently available at NPS in card or tape form, and considerable modification will be required to adapt them to use on the IBM 3033 system. In the absence of later forms of LOWTRAN, the IIIB version may still be used to compare transmittances under different conditions. In addition, little change has been made in the absorption in later versions, so that the use of IIIB to compute the atmospheric absorption is valid. This allows its use to compute aerosol extinction from total optical extinction by subtraction of the calculated absorption.

It is recommended that an effort should be made to adapt LOWTRAN V for the NPS IBM computer system.

APPENDIX A LOWTRAN IIIB PROGRAM

```

C*****
C*      NAVAL POSTGRADUATE SCHOOL MODIFICATION
C*      OF AFGL LOWTRAN 3B
C*      (1 OCTOBER 1976 VERSION)
C*
C*      MAJ. MOON-SIK SHIN
C*      AIR FORCE KOREA
C*
C*****
C*      PROGRAM LOWTRAN 3B CALCULATES THE TRANSMITTANCE OF THE ATMOS-
C*      PHERE FROM 350 CM-1 TO 40000 CM-1 (0.25 TO 28.57 MICRONS) AT
C*      20 CM-1 SPECTRAL RESOLUTION ON A LINEAR WAVELENGTH SCALE.
C*      REFRACTION AND EARTH-S CURVATURE EFFECTS ARE INCLUDED.
C*      THE ATMOSPHERE IS LAYERED IN ONE KILOMETER INTERVALS BETWEEN
C*      GROUND LEVEL AND 25 KM., 5 KM. INTERVALS TO 50 KM. A TWENTY
C*      KM. INTERVAL TO 70 KM., AND A THIRTY KM. INTERVAL TO 100 KM.
C*
C*****
C*      PROGRAM ACTIVATED BY SUBMISSION OF FOUR CARD SEQUENCE.
C*
C*****
C*      CARD 1: MODEL, IAERO, IHAZE, JP, IM, M1, M2, M3, ML, RO, VIS
C*      -----FORMAT(9I3, 3X, 2F10.3) *****
C*
C*      MODEL SELECTS ONE OF THE FOLLOWING MODEL ATMOSPHERE.
C*      MODEL = 0 FOR HORIZ. PATH WHEN METEOROL. DATA USED.
C*      INSTEAD OF CARD 2. READ H1, P(MB), T(DEG C), DEW PT. TEMP
C*      (DEG C), REL. HUMIDITY(%), H2O DENSITY(GM.M-3), O3 DENSITY
C*      (GM.M-3), RANGE(KM)
C*      -----FORMAT(3F10.3, 2F5.2, 2(1PE10.3), F10.3) .
C*      MODEL = 1 SPECIFIES A TROPICAL ATMOSPHERE.
C*      MODEL = 2 SPECIFIES A MIDLATITUDE SUMMER ATMOSPHERE.
C*      MODEL = 3 SPECIFIES A MIDLATITUDE WINTER ATMOSPHERE.
C*      MODEL = 4 SPECIFIES A SUB-ARCTIC SUMMER ATMOSPHERE.
C*      MODEL = 5 SPECIFIES A SUB-ARCTIC WINTER ATMOSPHERE.
C*      MODEL = 6 SPECIFIES A 1962 US STANDARD ATMOSPHERE.
C*      MODEL = 7 FOR NEW MODEL ATMOSPHERE (E.G. RADIOSCNDE DATA).
C*      READ BETWEEN CARDS 1 AND 2. ALTITUDE(KM), P(MB), T(DEG C),
C*      DEW PT. TEMP(DEG C), REL. HUMIDITY(%), H2O DENSITY(GM.M-3),
C*      O3 DENSITY(GM.M-3).
C*      -----FORMAT(3F10.3, 2F5.2, 2(1PE10.3), F10.3)
C*      NOTE EITHER DEW PT. TEMP, REL. HUM., OR H2O DENSITY
C*      CAN BE USED.
C*
C*      IAERO SELECTS THE TYPE OF AEROSOL ATTENUATION.
C*      IAERO = 1 SPECIFIES AN AVERAGE CONTINENTAL AEROSOL MODEL.

```

```

C* IAERO = 2 SPECIFIES A RURAL AEROSOL MODEL.
C* IAERO = 3 SPECIFIES AN URBAN AEROSOL MODEL.
C* IAERO = 4 SPECIFIES A MARITIME AEROSOL MODEL.
C*
C* IHAZE SELECTS THE DEGREE OF AEROSOL ATTENUATION.
C* IF IHAZE=0 NO AEROSOL SCATTERING IS COMPUTED.
C* IF IHAZE = 1 AND VIS IS NON-ZERO, THEN AEROSOL ATTENUATION
C* FOR THE VISIBLE RANGE IS USED.
C* IF IHAZE = 1 OR 2 AND VIS IS ZERO, THEN AEROSOL ATTENUATION
C* FOR 23 KM AND 5 KM VISIBLE RANGES, RESPECTIVELY,
C* IS USED.
C*
C* JP SELECTS THE PRINT OPTION.
C* JP = 0 FOR NORMAL OPERATION.
C* JP = 1 TO SUPPRESS THE PRINTING OF THE TRANSMITTANCE TABLE.
C*
C* IM DETERMINES WHEN RADIOSONDE DATA IS TO BE READ IN.
C* IM = 0 FOR NORMAL OPERATION.
C* IM = 1 FOR INITIALIZING RADIOSONDE OR METEOROLOGICAL
C* DATA.
C*
C* M1, M2, AND M3 ARE USED TO CHANGE TEMP, H2O, AND O3 ALTITUDE
C* PROFILES, RESPECTIVELY, TO ANOTHER MODEL VALUE.
C*
C* ML IS THE NUMBER OF LEVELS OF RADIOSONDE DATA FOR MODEL = 7.
C*
C* R0 IS THE INPUT VALUE FOR THE RADIUS OF THE EARTH. IF R0 IS
C* ZERO THEN THE PROGRAM USES STORED VALUES.
C*
C* VIS IS THE VISUAL RANGE AT SEA LEVEL (KM)
C*
C**** CARL 2: ITYPE, LEN, H1, H2, ANGLE, RANGE, BETA---FORMAT(2I3,4X,5F10.3)
C*
C* ITYPE INDICATES THE TYPE OF ATMOSPHERIC PATH.
C* ITYPE = 1, CORRESPONDS TO A HORIZONTAL (CONSTANT PRESSURE)
C* PATH. READ H1 AND RANGE.
C* ITYPE = 2, VERTICAL OR SLANT PATH BETWEEN TWO ALTITUDES.
C* READ H1 AND TWO OTHER GEOMETRIC PARAMETERS (E.G., H2 AND
C* ANGLE).
C* ITYPE = 3, VERTICAL OR SLANT PATH TO SPACE.
C* READ H1 AND ANGLE.
C*
C* LEN SELECTS THE TYPE OF RAY PATH TO BE USED.
C* LEN = 0 FOR NORMAL OPERATION OF THE PROGRAM WHICH SELECTS
C* THE SHORTER PATH WHEN APPLICABLE.
C* LEN = 1 TO SELECT THE LONGER PATH WHEN APPLICABLE.
C*

```

```

*LL 00470
*LL 00480
*LL 00490
*LL 00500
*LL 00510
*LL 00520
*LL 00530
*LL 00540
*LL 00550
*LL 00560
*LL 00570
*LL 00580
*LL 00590
*LL 00600
*LL 00610
*LL 00620
*LL 00630
*LL 00640
*LL 00650
*LL 00660
*LL 00670
*LL 00680
*LL 00690
*LL 00700
*LL 00710
*LL 00720
*LL 00730
*LL 00740
*LL 00750
*LL 00760
*LL 00770
*LL 00780
*LL 00790
*LL 00800
*LL 00810
*LL 00820
*LL 00830
*LL 00840
*LL 00850
*LL 00860
*LL 00870
*LL 00880
*LL 00890
*LL 00900
*LL 00910
*LL 00920
*LL 00930

```



```

C          INPUT SPECTRAL DATA FOR OZONE
C
C      DO 6 J=1,540,10
C      K=J+9
C      50 FORMAT(8X,10(1X,F5.2))
C      6 READ(5,50) (C3(I),I=J,K)
C
C      WRITE(6,69)
C      69 FORMAT(8X,1X,'SPECTRAL DATA OZONE',/)
C      WRITE(6,51)
C      51 FORMAT(7X,1X,'C301 DATA',/)
C      WRITE(6,50) (C3(I),I=1,190)
C
C      WRITE(6,52)
C      52 FORMAT(7X,1X,'C302 DATA',/)
C      WRITE(6,50) (C3(I),I=191,386)
C
C      WRITE(6,53)
C      53 FORMAT(7X,1X,'C303 DATA',/)
C      WRITE(6,50) (C3(I),I=387,540)
C
C      *** INPUT SPECTRAL DATA FOR N2 CONTINUUM
C      READ(5,72) (C4(I),I=1,5)
C      72 FORMAT(8X,5(1X,1PE8.2))
C
C      L=0
C      DO 7 J=1,18
C      L=L+6
C      M=I+6
C      READ(5,55) (C4(I),I=L,M)
C      55 FORMAT(8X,7(1X,F8.2))
C      L=L+1
C      7 CONTINUE
C
C      READ(5,73) (C4(I),I=132,133)
C      73 FORMAT(8X,2(1X,1PE8.2))
C
C      WRITE(6,54)
C      54 FORMAT(7X,1X,'C4 SPECTRAL DATA- N2 CONTINUUM',/)
C      WRITE(6,72) (C4(I),I=1,5)
C      WRITE(6,14) (C4(I),I=6,131)
C      14 FORMAT(1X,1X,1PE8.2)
C      WRITE(6,73) (C4(I),I=132,133)
C
C      *** SPECTRAL DATA H2O CCNTINUUM (4 MICRON REGION DATA)

```

```

LL 01480
LL 01490
LL 01500
LL 01510
LL 01520
LL 01530
LL 01540
LL 01550
LL 01560
LL 01570
LL 01580
LL 01590
LL 01600
LL 01610
LL 01620
LL 01630
LL 01640
LL 01650
LL 01660
LL 01670
LL 01680
LL 01690
LL 01700
LL 01710
LL 01720
LL 01730
LL 01740
LL 01750
LL 01760
LL 01770
LL 01780
LL 01790
LL 01800
LL 01810
LL 01820
LL 01830
LL 01840
LL 01850
LL 01860
LL 01870
LL 01880
LL 01890
LL 01900
LL 01910
LL 01920
LL 01930
LL 01940
LL 01950
LL 01960

```



```

C      READ(5,62) (C703(I), I=43,45)
C      WRITE(6,66)
C      66  FORMAT(8X,/, 1X, 'URBAN AEROSOL MODEL', /)
C      WRITE(6,61) (C703(I), I=1,45)
C      ** MARITIME AEROSOL MODEL
C      DO 67 J=1,42,7
C      K=J+6
C      67  READ(5,59) (C704(I), I=J,K)
C      READ(5,62) (C704(I), I=43,45)
C      WRITE(6,68)
C      68  FORMAT(8X,/, 1X, 'MARITIME AEROSOL MODEL', /)
C      WRITE(6,61) (C704(I), I=1,45)
C      ** SPECTRAL DATA AEROSOL ABSORPTION
C      ** AVERAGE CONTINENTAL AEROSOL MODEL
C      DO 16 J=1,42,7
C      K=J+6
C      16  READ(5,59) (C7A01(I), I=J,K)
C      READ(5,62) (C7A01(I), I=43,45)
C      WRITE(6,61)
C      17  FORMAT(8X,/, 1X, 'SPECTRAL DATA AEROSOL ABSORPTION *** AVERAGE
C      *CONTINENTAL AEROSOL MODEL', /)
C      WRITE(6,61) (C7A01(I), I=1,45)
C      ** RURAL AEROSOL MODEL
C      DO 70 J=1,42,7
C      K=J+6
C      70  READ(5,59) (C7A02(I), I=J,K)
C      READ(5,62) (C7A02(I), I=43,45)
C      WRITE(6,61)
C      71  FORMAT(8X,/, 1X, 'RURAL AEROSOL MODEL', /)
C      WRITE(6,61) (C7A02(I), I=1,45)
C      ** URBAN AEROSOL MODEL
C      DO 74 J=1,42,7
C      K=J+6
C      74  READ(5,59) (C7A03(I), I=J,K)
C      READ(5,62) (C7A03(I), I=43,45)
C      WRITE(6,75)

```

```

LL 02450
LL 02460
LL 02470
LL 02480
LL 02490
LL 02500
LL 02510
LL 02520
LL 02530
LL 02540
LL 02550
LL 02560
LL 02570
LL 02580
LL 02590
LL 02600
LL 02630
LL 02640
LL 02650
LL 02660
LL 02680
LL 02690
LL 02700
LL 02710
LL 02720
LL 02730
LL 02740
LL 02750
LL 02760
LL 02770
LL 02780
LL 02790
LL 02800
LL 02810
LL 02820
LL 02830
LL 02840
LL 02850
LL 02860
LL 02870
LL 02880
LL 02890
LL 02900
LL 02910
LL 02920
LL 02930
LL 02940

```

```

75 FORMAT(8X,'1X, URBAN AEROSOL MODEL ',/)
CC WRITE(6,61) (C7A03(I), I=1,45)
C
C *** MARITIME AEROSOL MODEL
C
DO 76 J=1,42,7
K=J+6
76 READ(5,59) (C7A04(I), I=J,K)
CC READ(5,62) (C7A04(I), I=43,45)
C WRITE(6,77)
77 FORMAT(8X,'1X, MARITIME AEROSOL MODEL ',/)
CC WRITE(6,61) (C7A04(I), I=1,45)
C
C *** SPECTRAL DATA - OZONE - UV AND VISIBLE
C
REAL(5,72) (C8(I), I=1,5)
L=0
DO 78 J=1,13
L=L+6
M=L+6
78 REAL(5,55) (C8(I), I=L,M)
C L=L+1
C CONTINUE
79 READ(5,79) (C8(I), I=97,102)
CC FORMAT(8X,6(1X,E8.2))
C
C WRITE(6,80)
C
80 FORMAT(8X,'1X, SPECTRAL DATA-- OZONE - UV AND VISIBLE',/)
CC WRITE(6,72) (C8(I), I=1,5)
C WRITE(6,14) (C8(I), I=6,102)
C
C *** INEUT PARTIAL AMT OF C2 DATA
C
DO 83 J=951,1570,10
K=J+9
83 READ(5,50) (C2(I), I=J,K)
CC READ(5,50) (C2(I), I=1571,1575)
C
C WRITE(6,84)
C
84 FORMAT(8X,'1X, SPECTRAL DATA*** UNIFORMLY MIXED GASES C206 DATA',/)
CC WRITE(6,50) (C2(I), I=951,1140)
C
C WRITE(6,85)
C
85 FORMAT(8X,'1X, C207 DATA',/)
CC WRITE(6,56) (C2(I), I=1141,1330)
C WRITE(6,81)
C

```

```

02950
LL 02960
LL 02970
LL 02980
LL 02990
LL 03000
LL 03010
LL 03020
LL 03030
LL 03040
LL 03050
LL 03060
LL 03070
LL 03100
LL 03110
LL 03120
LL 03130
LL 03140
LL 03150
LL 03160
LL 03170
LL 03180
LL 03190
LL 03200
LL 03210
LL 03220
LL 03230
LL 03240
LL 03250
LL 03260
LL 03280
LL 03290
LL 03310
LL 03320
LL 03330
LL 03340
LL 03350
LL 03360
LL 03370
LL 03380
LL 03390
LL 03400
LL 03410
LL 03420
LL 03430
LL 03440
LL 03450

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```

CC      81 FORMAT(8X,56,'1X,C208 DATA',1520)
CC      WRITE(6,56) (C2(I),I=1331,1520)
CC
CC      82 FORMAT(8X,56,'1X,C209 DATA',1575)
CC      WRITE(6,56) (C2(I),I=1521,1575)
CC
C*****
C BEGIN MAIN CODE
C*****
CALL ARRAY (P,T,WH,WO,VX,C1,C2,C7,C7A)
IXY=2
IP=1
KL1=NL-1
DO 108 ITER=1,10000
  IF (IXY.LE.0.OR.IXY.GE.5) GO TO 209
  IF (IXY.EQ.1) GO TO 107
  IF (IXY.EQ.3) GO TO 105
  READ(5,400) MODEL,IAERO,IHAZE,JP,IM,M1,M2,M3,ML,R0,VIS
  IF (MODEL.EQ.1) WRITE(6,410)
  IF (MODEL.EQ.2) WRITE(6,411)
  IF (MODEL.EQ.3) WRITE(6,412)
  IF (MODEL.EQ.4) WRITE(6,413)
  IF (MODEL.EQ.5) WRITE(6,414)
  IF (MODEL.EQ.6) WRITE(6,415)
  IF (M1.NE.0) WRITE(6,444) M1
  IF (M2.NE.0) WRITE(6,445) M2
  IF (M3.NE.0) WRITE(6,446) M3
  IF (IAERO.EQ.1) WRITE(6,447)
  IF (IAERO.EQ.2) WRITE(6,448)
  IF (IAERO.EQ.3) WRITE(6,449)
  IF (IAERO.EQ.4) WRITE(6,450)
  IF (IHAZE.EQ.0) WRITE(6,426)
  IF (IHAZE.EQ.1) AND:CKZERO(VIS) VIS=23.
  IF (IHAZE.EQ.2) AND:CKZERO(VIS) VIS=5.
  IF (.NOT.CKZERO(VIS)) WRITE(6,417) VIS
  IF (VIS.GT.0.0) AND:VIS.LT.2.0) WRITE(6,442)
  IF (MODEL.EQ.0) M=MODEL
  IF (MODEL.NE.0) NLP=1
  IF (MODEL.EQ.0) AND:MODEL.NE.7) NLP=NL
  IF (MODEL.NE.0) AND:MODEL.NE.7) NLP=MAX0(ML,1)
  IF (ML.GT.NL) WRITE(6,451) ML,NL
  IF (ML.GT.NL) STOP
DO 101 I = 1,NL
  IF (M.NE.7) AHZ1(I)=HZ1(I)

```

```

101 IF (M.NE.7) AHZ2(I)=HZ2(I)
    IF (M.NE.7) Z(I)=Z0(I)
    CONTINUE
    IF (CKZERO(R0)) RE=REARTH(M)
    IF (.NOT.CKZERO(R0)) RE=R0
    IF (M.NE.7.OR.IN.EQ.0) GO TO 104

C**** READ IN RADIOSCNDE (MODEL = 7) OR METEOROLOGICAL (MODEL = 0) DATA
C
100 CONTINUE
    DC 103 K = 1,NLP
    IF (MODEL.EQ.0) READ(5,429) Z(K),P(7,K),TMP,DP,RH,WH(7,K),
    *
    IF (MODEL.EQ.7) READ(5,429) Z(K),P(7,K),TMP,DP,RH,WH(7,K),
    *
    IF (M(7,K),AHZ1(K)=0.0
    IF (MODEL.EQ.0) H1=Z(K)
    DO 102 I = 1,NL1
    IF (Z(K).GE.Z0(I)) J=I
    FAC=(Z(K)-Z0(J))/Z0(J+1)-Z0(J))
    T(7,K)=TMP+273.15
    IF (M1.NE.0) T(7,K)=T(M1,J)+1/T(M1,J)) **FAC
    IF (.NOT.CKZERO(RH)) TT=273.15/(273.15+DP)
    IF (.NOT.CKZERO(DP)) TT=273.15/(273.15+DP)
    IF (CKZERO(WH(7,K))) WH(7,K)=EXP(18.9766-14.9595*TT-2.43882
    *
    IF (.NOT.CKZERO(RH)) WH(7,K)=0.01*RH*WH(7,K)
    IF (M2.NE.0) WH(7,K)=WH(M2,J)*WH(M2,J+1)/WH(M2,J)) **FAC
    IF (M3.NE.0) WH(7,K)=WH(M3,J)*WH(M3,J+1)/WH(M3,J)) **FAC
    IF (CKZERO(AHZ1(K))) AHZ2(K)=AHZ2(J+1)/AHZ2(J)) **FAC
    IF (.NOT.CKZERO(AHZ1(K))) AHZ2(K)=AHZ1(K)
    IF (.NOT.CKZERO(AHZ1(K))) AHZ1(K)=AHZ2(K)
    IF (CKZERO(AHZ1(K))) AHZ1(K)=AHZ2(K)
    IF (MODEL.EQ.0) WRITE(6,430) Z(K),P(7,K),TMP,DP,RH,WH(7,K),
    *
    IF (MODEL.EQ.7) WRITE(6,429) Z(K),P(7,K),TMP,DP,RH,WH(7,K),
    *
    CONTINUE
103 CCNTINUE
C
104 IF (M1.EQ.0) M1=M
    IF (M2.EQ.0) M2=M
    IF (M3.EQ.0) M3=M
    IF (MODEL.EQ.0) ITYPE=1
    IF (IXY.EQ.4) GO TO 108
    CONTINUE
105

```

```

C
IF (MODEL.EQ.0) GO TO 100
READ(5,406) ITYPE, LEN, H1, H2, ANGLE, RANGE, BETA
X1=RE+H1
IFIND=0
IF (CKZERO(RANGE) .AND. ((ITYPE.EQ.2 .AND. .NOT. CKZERO(BETA)))
  .OR. (ITYPE.EQ.3 .AND. H2.LT.H1 .AND. H2.GT.0.0)) IFIND=1
IF (ITYPE.EQ.1 .OR. ITYPE.EQ.3) GO TO 106
ITYPE=2
IF (.NOT. CKZERO(ANGLE) .AND. .NOT. CKZERO(BETA)) H2=
  X1*ABS((SIN(ANGLE*CA)/SIN((ANGLE-BETA)*CA)))-RE
IF (.NOT. CKZERO(ANGLE) .AND. .NOT. CKZERO(RANGE)) H2=
  X1*SQRT(1+((RANGE/X1)**2)*((RANGE/X1)*COS(ANGLE*CA))-RE
IF (.NOT. CKZERO(BETA) .AND. .NOT. CKZERO(RANGE)) H2=
  X1*ABS(COS(BETA*CA)+SQRT((RANGE/X1)**2-SIN(PETA*CA)**2))-RE
X2=RE+H2
IF (.NOT. CKZERO(RANGE)) ANGLE=ARCCOS(((X2/X1)*(X2/RANGE)-
  (X1/RANGE)-(RANGE/X1))/2.)/CA
IF (.NOT. CKZERO(BETA*CA)) ANGLE=ATAN(X2*SIN(BETA*CA)/(X2*
  COS(BETA*CA)-X1))/CA
IF (ANGLE.LT.0.0) ANGLE=ANGLE+180.
IF (CKZERO(RANGE)) RANGE=X1*ABS(SQRT((X2/X1)**2-SIN(ANGLE*CA)
  **2))-COS(ANGLE*CA)
IF (CKZERO(BETA)) BETA=ARSIN((RANGE/X2)*SIN(ANGLE*CA))/CA
CONTINUE
WRITE(6,428) H1, H2, ANGLE, RANGE, BETA
IF (ITYPE.EQ.1) WRITE(6,407) H1, RANGE
IF (ITYPE.EQ.2) WRITE(6,408) H1, H2, ANGLE
IF (ITYPE.EQ.3) WRITE(6,409) H1, ANGLE
IF (ITYPE.EQ.3) H2=Z(NLP)

C 107
CONTINUE
IF (IXY.EQ.3 .OR. IXY.EQ.4) GO TO 108
READ(5,405) V1, V2, DV
WL2=1000./V1
WL1=1000./V2
IF (IXY.LE.2) WRITE(6,418) V1, V2, DV, WL1, WL2
WL=SQRT(WL1*WL2)
CO=77.46+0.459/WL**2
CW=43.487-0.3473/WL**2
IF (IXY.EQ.1) GO TO 49

C 108
CONTINUE
IF (IFIND.EQ.1) CALL ANGL (H1, H2, ANGLE, BETA, LEN, NLP)
IF (IFIND.EQ.1) GO TO 110
IF (JP.EQ.0) WRITE(6,427)
IF (ITYPE.EQ.1) GO TO 1100
DO 109 K=1,10

```

```

LL 04500
LL 04510
LL 04520
LL 04530
LL 04540
LL 04550
LL 04560
LL 04570
LL 04580
LL 04590
LL 04600
LL 04610
LL 04620
LL 04630
LL 04640
LL 04650
LL 04660
LL 04670
LL 04680
LL 04690
LL 04700
LL 04710
LL 04720
LL 04730
LL 04740
LL 04750
LL 04760
LL 04770
LL 04780
LL 04790
LL 04800
LL 04810
LL 04820
LL 04830
LL 04840
LL 04850
LL 04860
LL 04870
LL 04880
LL 04890
LL 04900
LL 04910
LL 04920
LL 04930
LL 04940
LL 04950

```

```

109      VH(K)=0.0
        EETA=0.0
        SR=0.0
C***** NOW DEFINE CONSTANT PRESSURE PATH QUANTITIES EH(1-8)
C
        SPHI=SIN(ANGLE*CA)
        R1=(RE+H1)*SPHI
        IF (H1.LE.Z(NLP)) GO TO 110
        X=(RE+Z(NLP))/(RE+H1)
        IF (SPHI.GT.X) HMIN=R1-RE
        IF (SPHI.GT.X) WRITE(6,433) HMIN
        IF (SPHI.GT.X) GO TO 269
        H1=Z(NLP)
        J1=NLP
        SPHI=SPHI/X
        ANGLE=180.-ARSIN(SPHI)/CA
        CONTINUE
        R1=(RE+H1)*SPHI
        CONTINUE
        DO 112 I = 1,NLP
            PS=P(M,I)/1013.0
            TS=273.15/T(M1,I)
            TS1=(296.0/273.15)*TS
            X=PS*TS
            PT=PS*SORT(TS)
            D=0.1*WH(M2,I)
            PPW=4.56E-6*T(M1,I)*WH(M2,I)
C
C      *** THIS DEFINITION OF HAZE WHEN VIS = 0.0 MAY BE IN ERROR.
C
        IF (CKZEFO(VIS)) HAZE=0.0
        IF (.NOT.CKZERO(VIS)) HAZE=115.*((AHZ2(I)-AHZ1(I))/VIS+
            AHZ1(I)/5.0-AHZ2(I)/23.0)/18.
            EH(1,I)=L*PT**0.9
            EH(2,I)=X*PT**0.75
            EH(3,I)=46.667*WO(M3,I)*PT**0.4
            EH(4,I)=0.8*PT*X
            EH(5,I)=D*(PPW*EXP(6.08*(TS1-1.0))+0.002*(PS-PPW))
            EH(6,I)=X
            EH(7,I)=3.5336E-4*AMAX1(HAZE,0.0)
            EH(8,I)=46.667*WO(M3,I)
            EH(9,I)=0.0
            EH(10,I)=D*(0.12*PS+0.88*PPW)*EXP(4.56*(TS1-1.0))
            REF=CO*P(M,I)/T(M1,I)-4.56E-6*CW*WH(M2,I)*T(M1,I)
            IF (I.EQ.NLP) GO TO 111
            PPW=4.56E-6*T(M1,I+1)*WH(M2,I+1)

```

```

LL 04960
LL 04970
LL 04980
LL 04990
LL 05000
LL 05010
LL 05020
LL 05030
LL 05040
LL 05050
LL 05060
LL 05070
LL 05080
LL 05090
LL 05100
LL 05110
LL 05120
LL 05130
LL 05140
LL 05150
LL 05160
LL 05170
LL 05180
LL 05190
LL 05200
LL 05210
LL 05220
LL 05230
LL 05240
LL 05250
LL 05260
LL 05270
LL 05280
LL 05290
LL 05300
LL 05310
LL 05320
LL 05330
LL 05340
LL 05350
LL 05360
LL 05370
LL 05380
LL 05390
LL 05400
LL 05410
LL 05420

```

```

111      EH(9,I)=0.5E-6*(REF+CO*P(M,I+1)/T(N1,I+1)-PPW*CW)
        IP {IPIND.EQ.0. OR JP.EQ.0} WRITE(6,434) I,Z(I),
            {EH(K,I),K=1,10),REF
        CONTINUE
        IF (H1.GE.Z(I)) J1=I
        EH(9,I)=EH(9,I)+1.0
112      CONTINUE
        IF (MODEL.NE.0) CALL POINT (H1,YN1,J1,NP1,E,IP)
        TX1=E(9)
        IF (ITYPE.EQ.1) GO TO 47

C*** DOWNWARD TRAJECTORY
C
        IF (ANGLE.LE.90.) GO TO 19
        K2=0
        IF (NP1.EQ.1) J1=J1-1
        J2=J1+1
        JP1=J1+1
        IF ((H2.GT.Z(J1+1)) .OR. CKZERO(H2-H1)) .OR. (NP1.EQ.1.AND.
            {H2.GE.Z(J1+1)}) GO TO 30
        CALL POINT (H2,YN2,J2,NP2,W,IP)
        TX2=W(9)
        IF (H2.LT.H1) H=H2
        IF (J1.EQ.J2) TX2=TX1+YN2-EH(9,N)
        IF (H2.GT.H1) TX1=TX2
        IF (J1.EQ.J2.AND.H2.LT.H1) YN1=TX2
        A0=(H2+H1)*SEHI*YN1
        IF (H2.GE.H1) YN2=YN1
        DO 31 I=1,J1
            IF (I.NE.J1) HMIN=A0/EH(9,I)-RE
            IF (I.EQ.J1) HMIN=A0/YN1-RE
            IF (HMIN.LE.Z(I+1)) GO TO 32
        CONTINUE
        X=HMIN
        IF (HMIN.LE.0.00) GO TO 34

C*** THE USE OF YN BELCW MAY POSSIBLY BE IN ERROR
C      SHOULD BE YN1 OR YN2
C
        CALL POINT (HMIN,YN,JMIN,NP,TX,IP)
        TX3=TX(9)
        IF (J2.EQ.JMIN.OR.J1.EQ.JMIN) TX3=YN2+TX(9)-EH(9,N)
        IF (J1.EQ.JMIN.AND.H2.GE.H1) GO TO 33
        HMIN=A0/TX3-RE
        IF (ABS(X-HMIN).GT.0.0001) GO TO 32
        IF (J1.EQ.JMIN.AND.H2.GE.H1) YN1=TX3
        IF (J2.EQ.JMIN.AND.J1.NE.J2) YN2=TX3
33

```

```

LL 05430
LL 05440
LL 05450
LL 05460
LL 05470
LL 05480
LL 05490
LL 05500
LL 05510
LL 05520
LL 05530
LL 05540
LL 05550
LL 05560
LL 05570
LL 05580
LL 05590
LL 05600
LL 05610
LL 05620
LL 05630
LL 05640
LL 05650
LL 05660
LL 05670
LL 05680
LL 05690
LL 05700
LL 05710
LL 05720
LL 05730
LL 05740
LL 05750
LL 05760
LL 05770
LL 05780
LL 05790
LL 05800
LL 05810
LL 05820
LL 05830
LL 05840
LL 05850
LL 05860
LL 05870
LL 05880
LL 05890

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```

LL 05900
LL 05910
LL 05920
LL 05930
LL 05940
LL 05950
LL 05960
LL 05970
LL 05980
LL 05990
LL 06000
LL 06010
LL 06020
LL 06030
LL 06040
LL 06050
LL 06060
LL 06070
LL 06080
LL 06090
LL 06100
LL 06110
LL 06120
LL 06130
LL 06140
LL 06150
LL 06160
LL 06170
LL 06180
LL 06190
LL 06200
LL 06210
LL 06220
LL 06230
LL 06240
LL 06250
LL 06260
LL 06270
LL 06280
LL 06290
LL 06300
LL 06310
LL 06320
LL 06330
LL 06340
LL 06350
LL 06360

```

```

C 34
IF (H2.GE.H1) TX2=TX3
IF (H2.GE.H1) J2=JMIN H=HMIN
IF (H2.GE.H1.OR.H2.LT.HMIN) H=HMIN
WRITE(6,436) HMIN
IF (H2.LT.HMIN) WRITE(6,440) HMIN
GO TO 35

WRITE(6,436) HMIN
IF (H2.LT.H1) GO TO 35
IF (ITYPE.EQ.3.OR.H2.GE.H1) WRITE(6,437)
ITYPE=2
TX2=EH(9,1)
JMIN=0
J2=1
H2=0.0
H=0.0

C**** NOW DEFINE VERTICAL PATH QUANTITIES VH(1-8)
C 35
IF (JP.EQ.0) WRITE(6,420)
DO 135 ITES = 1,10000
IF (K2.EQ.0) REF=YN1
IF (K2.EQ.1) REF=YN2
X1=H1
J=JP1+1
CONTINUE
J=J-1
X2=Z(J)
IF (J.EQ.J2.AND.K2.EQ.0) X2=H
IF (J.EQ.JMIN.AND.K2.EQ.1) X2=HMIN
HN = (RE+X1)*SPHI-RE
IF (HM.GT.Z(J).AND.HM.GT.X2) X2=HM
RX = (RE+X1)/(RE+X2)
DS=X1-X2
ALP=90.0
THET=ARCSIN(SPHI)/CA
SALP=RX*SPHI
IF (ABS(X2-HM).GT.1.E-5) ALP=ARSIN(SALP)/CA
BET=ALP-THET
IF (SPHI.GT.1.E-10) DS=(RE+X2)*SIN(BET*CA)/SPHI
THETA=180.-THET
BETA=BETA+BET
ESI=BETA-ALP-ANGLE+180.
SR=SR+DS
DO 39 K = 1,10
AJ=EH(K,J)
BJ=EH(K,J+1)

```

```

LL 06370
LL 06380
LL 06390
LL 06400
LL 06410
LL 06420
LL 06430
LL 06440
LL 06450
LL 06460
LL 06470
LL 06480
LL 06490
LL 06500
LL 06510
LL 06520
LL 06530
LL 06540
LL 06550
LL 06560
LL 06570
LL 06580
LL 06590
LL 06600
LL 06610
LL 06620
LL 06630
LL 06640
LL 06650
LL 06660
LL 06670
LL 06680
LL 06690
LL 06700
LL 06710
LL 06720
LL 06730
LL 06740
LL 06750
LL 06760
LL 06770
LL 06780
LL 06790
LL 06800
LL 06810
LL 06820
LL 06830

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```

IP (J.EQ.J1) BJ=E(K)
IF (J.EQ.J2) AND.H2.LT.H1.AND.H2.GT.0.0) AJ=W(K)
IF (J.EQ.JMIN.AND.H2.GE.H1) AJ=TX(K)
IF (J.EQ.JMIN.AND.ABS(H2-H1).LT.1.E-5) AJ=TX(K)
IF (K2.NE.0.AND.J.EQ.J2) BJ=W(K)
IF (K2.NE.0.AND.J.EQ.JMIN) AJ=TX(K)
IF (CKZERO(AJ).OR.CKZERO(BJ)) EV=0.0
IF (CKZERO(AJ-BJ)) EV=DS*AJ
IF (.NOT.CKZERO(AJ).AND.NOT.CKZERO(BJ).AND.NCT.CKZERO(AJ-BJ)) EV=DS*(AJ-BJ)/ALOG(AJ/BJ)
VH(K)=VH(K)+EV
CONTINUE
IF (JP.EQ.0) WRITE(6,435) J,X1,(VH(L),L=1,8),PSI,ALP,
    BETA,THETA,SR
IF (J.EQ.J2.AND.H2.GE.H1) GO TO 45
IF (J.EQ.JMIN.AND.K2.EQ.1) GO TO 43
IF (J.NE.1) RN=REF/TH(9,J-1)
IF (J.EQ.J2+1) RN=REF/TH2
IF (J.EQ.J2.AND.K2.EQ.0) RN=REF/YN2
IF (J.EQ.(JMIN+1).AND.K2.EQ.1) RN=REF/TH3
IF (SALP.GE.RN) RN=1.0
SPHI=SALP*RN
REF=TH(9,J)
IF (J.EQ.J2+1.AND.K2.EQ.0) REF=TX2
X1=X2
IF (J.NE.1.AND.(J.NE.J2.OR.K2.NE.0)) GO TO 38
IF (HMIN.LE.0.0) GO TO 47
IF (LEN.EQ.0) WRITE(6,438)
IF (LEN.EQ.0) GO TO 47
WRITE(6,439)
K2=1
IF (ABS(X1-HMIN).LE.0.001) GO TO 47
H=HMIN
JP1=J2+1
IF (NP2.EQ.1) JP1=JP1-1
B=BETA
PH=180.-ARCSIN(SPHI)/CA
PS=SR
DO 42 K=1,10
    E(K)=VH(K)
CONTINUE
BETA=2.*BETA-B
PSI=2.*PSI-PS
SR=2.*SR-TS
C**** LCNG PATH TAKEN

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C
44      PHI=PH
44      DO 44 K = 1, 10
44      VH(K)=2.*VH(K)-E(K)
44      GO TO 47
C
45      CCNTINUE
45      DO 46 K = 1, 10
46      VH(K)=2.*VH(K)
46      EETA=2.*BETA
46      SR=2.*SR
46      IF (CKZERO (H2-H1)) GO TO 47
46      RN=TX1/YN1
46      SPHI=SIN(ANGLE*CA)
46      IF (SPHI.LT.EN) SPHI=SPHI/RN
C
C****
C      UPWARD TRAJECTORY
C      19      IP (ANGLE.GT.90..AND.NP1.GT.0) J1=J1+1
C      IF (ITYPE.EQ.3) J2=NLP
C      IF (ITYPE.EQ.2) CALL POINT(H2,YN2,J2,NP,TX,IP)
C      IF (ITYPE.EQ.2) AND.NP.EQ.1) J2=J2-1
C      EH(10,J1)=E(10)
C      DO 21 K = 1, 10
C      IF (ITYPE.EQ.3) EH(K,J1)=E(K)
C      IF (ITYPE.EQ.2) EH(K,J2+1)=TX(K)
C      CONTINUE
C      21      IF (ITYPE.EQ.2) EH(10,J2+1)=TX(10)
C      IF (J1.EQ.J2) TX1=TX1+YN2-EH(9,J1)
C
C****
C      NOW DEFINE VERTICAL PATH QUANTITIES VH(1-8)
C      IF (JP.EQ.0) WRITE(6,420)
C      X1=H1
C      DO 25 I = J1,J2
C      X2=Z(I+1)
C      IF (I.EQ.J2) X2=H2
C      DZ=X2-X1
C      IF (I.EQ.NLP) DZ=Z(I)-Z(I-1)
C      RX={RE+X1}/{RE+X2}
C      THETA=ARCSIN(SPHI)/CA
C      PHI=ARCSIN(SPHI*RX)/CA
C      BET=THETA-PHI
C      SALP=RX*SPHI
C      IF (SPHI.GT.1.E-10) DZ=(RE+X2)*SIN(BET*CA)/SPHI
C      BETA=BETA+BET
C      PSI=BETA+PHI-ANGLE

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LL 06840
LL 06850
LL 06860
LL 06870
LL 06880
LL 06890
LL 06900
LL 06910
LL 06920
LL 06930
LL 06940
LL 06950
LL 06960
LL 06970
LL 06980
LL 06990
LL 07000
LL 07010
LL 07020
LL 07030
LL 07040
LL 07050
LL 07060
LL 07070
LL 07080
LL 07090
LL 07100
LL 07110
LL 07120
LL 07130
LL 07140
LL 07150
LL 07160
LL 07170
LL 07180
LL 07190
LL 07200
LL 07210
LL 07220
LL 07230
LL 07240
LL 07250
LL 07260
LL 07270
LL 07280
LL 07290
LL 07300

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PHI=180.-EHI
SR=SR+DZ
IO=MAXO(MINO(I,NLP-1),1)
DO 24 K=1,IO
  IF (CKZERO(EH(K,IO)).OR. CKZERO(EH(K,IO+1)).OR.
    EV=DZ*AMIN1(EH(K,IO),EH(K,IO+1))
  IF (.NOT. CKZERO(EH(K,IO)).AND..NOT. CKZERO(EH(K,IO+1))
    .AND..NOT. CKZERO(EH(K,IO)-EH(K,IO+1)).AND.I.NE.NLP)
    EV=DZ*(EH(K,IO)-EH(K,IO+1))/ALOG(EH(K,IO)/
    EH(K,IO+1))
  IF (.NOT. CKZERO(EH(K,IO)).AND..NOT. CKZERO(EH(K,IO+1))
    .AND..NOT. CKZERO(EH(K,IO)-EH(K,IO+1)).AND.I.EQ.NLP)
    EV=DZ*EH(K,IO+1)/ALOG(EH(K,IO)/EH(K,IO+1))
  VH(K)=VH(K)+EV
CONTINUE
IF (JP.EQ.0) WRITE(6,435) I,X1,(VH(L),L=1,8),PSI,PHI,BETA,
  THETA,SR
  IF (I.EQ.NLP) GO TO 25
  IF (I.EQ.J2-1) EH(9,I+1)=YN2
  IF (I.EQ.J1) EH(9,I)=TX1
  RN=EH(9,I+1)/EH(9,I)
  SPHI=SPHI+RX/RN
  IF (SALP.GE.RN) SPHI=SALP
  X1=X2
CONTINUE
CONTINUE
IF (ITYPE.NE.1) WRITE(6,443) HM
DO 48 K=1,IO
  IF (ITYPE.EQ.1.AND.MODEL.NE.0) W(K)=RANGE*E(K)
  IF (ITYPE.EQ.1.AND.MODEL.EQ.0) W(K)=RANGE*EH(K,1)
  IF (ITYPE.NE.1) W(K)=VH(K)
CONTINUE
WRITE(6,419)
WRITE(6,421) (W(I),I=1,8),W(10)
C**** BEGINNING OF TRANSMITTANCE CALCULATIONS
C
IV1=MAXO(5*IFIX(V1/5.0),350)
IV2=MINO(5*IFIX(V2/5.0+6.99),50000)
EV=AMAX1(DV,5.0)
IDV=IFIX(DV)
SUM=0.0
DO 206 IV=IV1,IV2,IDV
  IF (IV.GE.350.AND.IV.LT.9875) I={IV-350}/5+1
  IF (IV.GE.9875.AND.IV.LT.12800) I={IV-9875}/5+1771
  IF (IV.GE.12800.AND.IV.LT.13400) I={IV-12800}/5+2491

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C***** NITROGEN CONTINUUM
K=(IV-2080)/5+
IF {IV.LI.2080;OR.IV.GE.2740} TX(4)=0.0
IF {IV.GE.2080.AND.IV.LT.2740} TX(4)=C4(K)*W(4)

C***** WATER VAPOR CONTINUUM
XI=FLOAT(IV-2350)/50.0
NH=FIX(XI)
TX(5)=0.0

C*****
10 MICECN REGION
IF (IV.GE.670.AND.IV.LE.1350) TX(5)=(4.18+5578.0*
*EXP(-7.87E-3*FLOAT(IV)))*W(5)

C*****
4 MICECN REGION
IF (IV.GE.2350.AND.IV.LE.3000) TX(5)=(C5(NH)+ (XI-FLOAT(NH))
*(C5(NH+1)-C5(NH)))*W(10)

C***** MOLECULAR SCATTERING
IF (IV.LI.2740) TX(6)=0.0
IF {IV.GE.2740.AND.IV.LE.50000}
TX(6)=9.807E-20*(FLOAT(IV)**4.0117)*W(6)

C***** AEROSOL EXTINCTION
WL=10000./FLOAT(IV)
XX=0.0
YY=0.0
TX(7)=0.0
TX(10)=0.0
IF (IHAZE.EQ.0) GO TO 204
DO 203 L=1,44
IF (WL.GE.VX(IAERO,L)) NS=L
IF (NS.GT.0.AND.NS.LT.44) XX=C7(IAERO,NS+1)+
(C7(IAERO,NS)-C7(IAERO,NS+1))*(VX(IAERO,NS+1)-WL)/
(VX(IAERO,NS+1)-VX(IAERO,NS))
IF (NS.GT.0.AND.NS.LT.44) YY=C7A(IAERO,NS+1)+
(C7A(IAERO,NS)-C7A(IAERO,NS+1))*(VX(IAERO,NS+1)-WL)/
(VX(IAERO,NS+1)-VX(IAERO,NS))
CONTINUE
TX(7)=XX*W(7)
TX(10)=YY*W(7)

203
*
*
*
*
204

C***** UV OZONE
IF (IV.GE.13000.AND.IV.LE.23400) XI=FLOAT(IV-13000)/200.+1-LL
IF {IV.GE.27500.AND.IV.LE.50000} XI=FLOAT(IV-27500)/500.+57-LL

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```

LL 08250
LL 08260
LL 08280
LL 08290
LL 08300
LL 08310
LL 08320
LL 08340
LL 08350
LL 08360
LL 08370
LL 08380
LL 08400
LL 08410
LL 08420
LL 08430
LL 08450
LL 08460
LL 08470
LL 08480
LL 08500
LL 08510
LL 08520
LL 08530
LL 08540
LL 08560
LL 08570
LL 08580
LL 08590
LL 08600
LL 08610
LL 08620
LL 08630
LL 08640
LL 08650
LL 08660
LL 08670
LL 08680
LL 08690
LL 08700
LL 08710
LL 08720
LL 08730
LL 08740
LL 08750
LL 08770
LL 08780

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[illegible]

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413 *FORMAT('1'/'10X,50H SUP-ARCTIC (60 DEG. LAT.) SUMMER MODEL ATMOSPHELL 09260
414 *FORMAT('1'/'10X,50H SUB-ARCTIC (60 DEG. LAT.) WINTER MODEL ATMOSPHELL 09270
415 *FORMAT('1'/'10X,36H 1962 U.S. STANDARD ATMOSPHERE MODEL) 09280
417 *FORMAT('1'/'10X,14H HAZE MODEL = ,F5.1,29H KM VISUAL RANGE AT SEA LEVEL 09290
418 *FORMAT('1'/'10X,22H FREQUENCY RANGE V1 = F7.1,14H CM-1 TO V2 = F7.1 09300
419 *FORMAT('1'/'10X,38H EQUIVALENT SEA LEVEL ABSORBER AMOUNTS//21X,110H 09310
420 *FORMAT('1'/'10X,13H THEFA 09320
421 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09330
422 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09340
423 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09350
424 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09360
425 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09370
426 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09380
427 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09390
428 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09400
429 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09410
430 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09420
431 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09430
432 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09440
433 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09450
434 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09460
435 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09470
436 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09480
437 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09490
438 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09500
439 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09510
440 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09520
441 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09530
442 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09540
443 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09550
444 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09560
445 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09570
446 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09580
447 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09590
448 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09600
449 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09610
450 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09620
451 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09630
452 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09640
453 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09650
454 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09660
455 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09670
456 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09680
457 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09690
458 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09700
459 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09710
460 *FORMAT('1'/'10X,10H W(1-8) = 8(1PE14.3)/74X,1PE14.3/ 09720

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442 *FORMAT(61H IF SO CONDITIONS MAY EXIST AT SEA LEVEL FOR THIS VISUAL
*BY THE TRANSMITTANCE AT 0.55 MICRONS)
443 *FORMAT(15X,29H ESTIMATED TANGENT ALTITUDE = F10.3,3H KM/(
444 *FORMAT(15X,41H TEMPERATURE PROFILE FOR MODEL ATMOSPHERE ;15}
445 *FORMAT(15X,41H WATER VAPOR PROFILE FOR MODEL ATMOSPHERE ;15}
446 *FORMAT(15X,41H OZONE PROFILE FOR MODEL ATMOSPHERE ;15}
447 *FORMAT(10X,26H CONTINENTAL AEROSOL MODEL)
448 *FORMAT(10X,20H RURAL AEROSOL MODEL)
449 *FORMAT(10X,20H URBAN AEROSOL MODEL)
450 *FORMAT(10X,23H MARITIME AEROSOL MODEL)
451 *FORMAT(10H1 *** ,15,37H INPUT LEVELS EXCEEDS MODEL LIMIT OF ,
*I5)
END
SUBROUTINE POINT (X,Y,N,NP,TX,IP)
C*****
C*
C* SUBROUTINE POINT COMPUTES THE MEAN REFRACTIVE INDEX ABOVE AND
C* BELOW A GIVEN ALTITUDE AND INTERPOLATES EXPONENTIALLY TO DETER-
C* MINE THE EQUIVALENT ABSORBER AMOUNTS AT THAT ALTITUDE.
C*
C*****
C*
C* X IS THE GIVEN ALTITUDE IN KM.
C* TX(9) AND YN ARE THE MEAN REFRACTIVITY ABOVE AND BELOW X,
C* RESPECTIVELY.
C* N IS THE INTEGRAL CORRESPONDING TO THE ATMOSPHERIC LEVEL AT
C* CR BELOW X.
C* IF X COINCIDES WITH THE ATMOSPHERE LEVEL, NP = 1,
C* OTHERWISE NP = 0.
C* TX(1-8) ARE ABSORBER AMOUNTS PER KM AT X.
C* IP IS AN INDICATOR; IP = 0 IMPLIES A CALCULATION OF REFRACTIVE
C* INDEX ONLY. OTHERWISE, THE EQUIVALENT ABSORBER AMOUNTS ARE
C* ALSO CALCULATED.
C*****
C*
C* CCMON /LOWTRN/ IATH NL, HZ1(34), HZ2(34), Z0(34), P01(34), P02(34),
C* P03(34), P04(34), P05(34), P06(34), T01(34), T02(34), T03(34), T04(34),
C* T05(34), T06(34), WH01(34), WH02(34), WH03(34), WH04(34), WH05(34),
C* WH06(34), W01(34), W02(34), W03(34), W04(34), W05(34), W06(34),
C* TR(67), F0(67), C101(190), C102(190), C103(190), C104(190),
C* C105(190), C106(190), C107(190), C108(190), C109(190), C110(190),
C* C111(190), C112(190), C113(190), C114(110), C201(190), C202(190),
C* C203(190), C204(190), C205(190), C3(540), C4(133), C5(15), C701(45),
C* C702(45), C703(45), C704(45), C7A01(45), C7A02(45), C7A03(45),

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C      *C7A04(45),C8(102),VX01(45),VX02(45),VX03(45),VX04(45),
      *EH(10,34),M,M1,M2,M3,RE,CW,CO,PI,CA,REARTH(7),Z(34),
      *DIMENSION TX(10),P(7,34),T(7,34),WH(7,34),WO(7,34),VX(4,45),
      *C7(4,45),C7A(4,45),C1(2580),C2(1575),
      *LOGICAL CKZERO
      REF(X,Y,Z)=CO*X/Y-4.56E-6*Z*Y*CW
      CKZERO(X)=ABS(X).LT.1.E-20
      CALL ARRAY(P,T,WH,WO,VX,C1,C2,C7,C7A)
      N=0
      X=AMAX1(X,0.0)
      N1=NI-1
      DO 101 I=1,N11
      IF (X.GT.Z(I)) N=I
      IF (.NOT. CKZERO(X-Z(N+1))) NP=0
      IF (N.EQ.0.OR.NE.EQ.1) GO TO 102
      IF (N.EQ.1) NP=1
      FAC=AHN1((X-Z(N+1))/(Z(N+1)-Z(N)))**FAC
      PX1=T(M,N)*P(M,N+1)/T(M,N)**FAC
      WX1=WH(M,N)*WH(M,N+1)/WH(M,N)**FAC
      TX(9)=(REF(P(M,N),T(M,N),WH(M,N))+REF(PX1,TX1,WX1))/2.
      YN=(REF(P(M,N),T(M,N),WH(M,N))+REF(PX1,TX1,WX1))/2.
      GO TO 103
101
C      CONTINUE
      TX(9)=(EH(9,N+1)-1.)*1.E+06
      IF (N.EQ.0) YN=0.0
      IF (N.NE.0) YN=(EH(9,N)-1.)*1.E+06
      CONTINUE
103
C      IF (IP.EQ.0) GO TO 105
      DO 104 K=1,10
      IF (N.EQ.0.OR. CKZERO(EH(K,N))) .AND. K.NE.9) TX(K)=0.0
      IF (N.NE.0 .AND. .NOT. CKZERO(EH(K,N))) .AND. K.NE.9) TX(K)=
      * EH(K,N)*(EH(K,N+1)/EH(K,N))**FAC
      CONTINUE
104
      WRITE(6,1) X,N,NE,TX(9),YN,(TX(K),K=1,8)
      CCCONTINUE
105
      TX(9)=1.+TX(9)*1.E-06
      YN=1.+YN*1.E-06
      RETURN
1
      FORMAT(10X,'FROM POINT - HEIGHT =',F10.4,9H KM,N=,I3,7H, NP
      * =I2,35H REFRACTIVITY ABOVE AND BELOW X =,2(1PE10.3)/10X,
      * 40H EQUIVALENT ABSORBER AMOUNTS PER KM AT X//20X,8(1PE10.3))

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11130
11140
11150
11160
11170
11180
11190
11200
11210
11220
11230
11240
11250
11260
11270
11280
11290
11300
11310
11320
11330
11340
11350
11360
11370
11380
11390
11400
11410
11420
11430
11440
11450
11460
11470
11480
11490
11500
11510
11520
11530
11540
11550
11560
11570
11580
11590

```

```

TN=THET
TM=IN-0.5*CA
DO 10 I TER = 1,10
  ANGLE=THET
  FBT=0.0
  B1=0.0
  BET1=0.0
  BET2=0.0
  FBT1=0.0
  FBT2=0.0
  FBT3=0.0
  IF (BETA.LE.0.0) GO TO 2
  IF (2.*THET-PI.GT.1.E-8) GO TO 9
  IF (IP.EQ.100) GO TO 6
  XMIN=X2*COS(BETA)-RE
  IF (XMIN.LT.H1) GO TO 8
  IF (XMIN.GE.H1) GO TO 4
  HM IN=H2
  H2=H1
  H1=HM IN
  ANGLE=0.5*PI
  THET=ANGLE
  SPHI=1.0
  ANG=ANGLE/CA
  IP=100
  CALL POINT (H1,YN,J1,NE,TX,IP)
  TX1=TX(9)
  CALL POINT (H2,YN,J2,NE,TX,IP)
  IF (NP.EQ.1) J2=J2-1
  IF (J1.EQ.J2) TX1=TX1+YN-EH(9,J1)
  CCNTINUE
  X1=RE+H1
  FBT=-TAN(THET)
  DO 7 J = J1,J2
    IF (J.NE.J2) X2=RE+Z(J+1)
    IF (J.EQ.J2) X2=RE+H2
    SALP=X1*SPHI/X2
    ALP=AR SIN(SALP)
    RN=EH(9,J+1)/EH(9,J)
    IF ((J+1).EQ.J2) RN=YN/EH(9,J)
    IF ((J+1).EQ.J1) RN=EH(9,J+1)/TX1
    IF ((J+1).EQ.J2.AND.J.EQ.J1) RN=YN/TX1
    FBT=FBT+IAN(THET)-TAN(ALP)
    E1=B1+THET-ALP
    TH1=THET/CA
    C=ALP/CA
    IF (CKZEFO(X2-(RE+H2))) C=PI-ALP

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2 3 4 5 6

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LL 11600
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LL 11620
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LL 11670
LL 11680
LL 11690
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LL 11900
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LL 11930
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LL 11960
LL 11970
LL 11980
LL 11990
LL 12000
LL 12010
LL 12020
LL 12030
LL 12040
LL 12050
LL 12060

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IF (SALP.GE.RN) RN=1.
SPHI=SALF/RN
THET=ARSIN(SPHI)
X1=X2
CONTINUE
IF (BETA.LE.0.0) THET=C
GO TO 26
CONTINUE
TANG=-TANG
ANGLE=PI-ANGLE
ANG=ANGLE/CA
IF (H1.LE.0.0) GO TO 3
CONTINUE
IP=101
CALL POINT (H1,YN1,J1,NP1,TX,IP)
TX1=TX(9)
IF (NP1.EQ.1) J1=J1-1
J2=ML
IF (H2.GE.H1) GO TO 13
CALL POINT (H2,YN2,J2,NP2,TX,IP)
TX2=TX(9)
IF (J1.EQ.J2) TX2=YN1+TX(9)-EH(9,J1)
J2F=J2-1
EO 102 J = J1,J2P
X1=RE+Z(J+1)
X2=RE+Z(J)
IF (J.EQ.J1) X1=RE+H1
IF (J.EQ.J2) X2=RE+H2
SALP=X1*SPHI/X2
HMIN=X1*SPHI-RE
IF (SALP.LE.1.0) GO TO 11
SALP=SPHI
IF (HMIN.GT.H2) GO TO 18
ALP=ARSIN(SALP)
THET=ARSIN(SPHI)
BET1=ALP-THET
BET1=BET1+BET
PB=TAN(ALP)
PB=PB-TAN(THET)
IF (J.NE.J1)
PB T1=FB T1+PB
TH1=THET/CA
BE=BET/CA
AL=ALP/CA
IF (CKZERO(X2-(RE+H2))) C=PI-ALP
REF=EH(9,J)

```

7

8

9

11

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102 IF (J.EQ.J1) REF=YN1
12 IF (J.EQ.J2) REF=TX2
13 IF (J.EQ.1) GO TO 12
   RN=EH(9,J)/EH(9,J-1)
   IF (J.EQ.J1) RN=YN1/EH(9,J-1)
   IF (J.EQ.J2) RN=REF/TX2
   IF (J.EQ.J2) RN=REF/YN2
   IF (SALP*GE.RN) RN=1.
   SPHI=SALP*RN
   IF (Z(J).IE.H2) GO TO 12
   CONTINUE
14 X1=X2
   IF (ABS (Z (J) -H2) .GE. 1.0E-10.OR.J.EQ.1) GO TO 14
   J=J-1
   X1=RE+Z (J+1)
   IF (J.EQ.J1) X1=RE+H1
   IF (J.EQ.J2.AND.J.NE.J1) X1=RE+H2
   X2=RE+Z (J)
   HMIN=X1*SPHI-RE
   IF (HMIN.LE.0.0) B1=BET1
   IF (HMIN.LE.0.0) LEN=0
   IF (HMIN.LE.0.0) FBET=FBT1
   IF (HMIN.LE.0.0) GO TO 26
   IF (Z(J).LT.HMIN) GO TO 18
   REF=EH(9,J)
   IF (J.EQ.J2) REF=YN
   SALP=X1*SPHI/X2
   ALF=ARSIN(SALP)
   THET=ARSIN(SPHI)
   FEET=ALP-THET
   FB=TAN(ALP)-TAN(THET)
   FBET2=FBT2+FE
   BET2=BET1+BET2
   EMIN=BET1+BET2
   AL=ALP/CA
   TH1=THET/CA
   RN=REF/EH(9,J-1)
   IF (SALP*GE.RN) RN=1.0
   SPHI=SALP*RN
   GO TO 13
17 TX3=YN1+TX(9)-EH(9,J1)
   YN1=TX3
   IF (ABS (H2-Z (J+1)) .LE. 1.0E-5) YN1=TX(9)
   IF (ABS (H1-Z (J+1)) .LE. 1.0E-5) YN1=TX(9)
   FN=1.0
   GO TO 19
18 CALL POINT (HMIN,YN,J2,NP,TX,IP)

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LL 12540
LL 12550
LL 12560
LL 12570
LL 12580
LL 12590
LL 12600
LL 12610
LL 12620
LL 12630
LL 12640
LL 12650
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LL 12670
LL 12680
LL 12690
LL 12700
LL 12710
LL 12720
LL 12730
LL 12740
LL 12750
LL 12760
LL 12770
LL 12780
LL 12790
LL 12800
LL 12810
LL 12820
LL 12830
LL 12840
LL 12850
LL 12860
LL 12870
LL 12880
LL 12890
LL 12900
LL 12910
LL 12920
LL 12930
LL 12940
LL 12950
LL 12960
LL 12970
LL 12980
LL 12990
LL 13000

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IP=102
TX3=TX (9)
IF (J.EQ.J1.AND.H2.GE.H1) GO TO 17
IF (J.EQ.J1.OR.J.EQ.J2) TX3=YN2+TX (9)-EH (9,J)
IF (HMIN.GT.H2) TX3=TX (9)
IF (J.EQ.J1.AND.HMIN.GT.H2) GO TO 17
RN=REP/TX3
IF (SALP.GE.FN) RN=1.
SPHI=SALP*RN
X=X1*SPHI-RI
DIF=ABS (HMIN-X)
HMIN=X
IF (DIF.GT.1.0E-5) GO TO 18
Y2=RE+HMIN
THET=ARCSIN (SPHI)
IF (CKZERO (RN)) FBT3=-TAN (THET)
IF (CKZERO (RN)) GO TO 20
DNX=(TX3-1.0)*ALOG ((TX3-1.0)/(REF-1.0))/(X2-X1)
FBT3=-TAN (THET)*(1.0-1.0/(1.0+TX3/(X2*DNX)))
BET=0.5*PI-THET
BET2=BET+BT
BMIN=BET1+BT2
IF (H2.GE.H1) GO TO 23
FET=BET1+2.*BT2
DB1=BETA-BET1
CB2=BET-BETA
CB3=ABS (BMIN-BETA)
IF ((DB3.GT.DB1.OR.DB3.GE.DB2).AND.DB2.GT.DE1) B1=BET1
IF ((DB3.GT.DB1.OR.DB3.GE.DB2).AND.DB2.GT.DB1) LEN=0
IF (DB3.GT.DB1.OR.DB3.GE.DB2).AND.DB2.GT.DB1 FBT=FBT1
IF (DB3.LE.DB1.OR.DB2.LE.DB1).AND.DB3.LE.DE1 B1=BET1+BT2
IF (DB3.LE.DB1.OR.DB2.LE.DB1).AND.DB3.LE.DE1 FBT=FBT1+FBT2+
FBT3
IF (DB3.GE.LE2.AND.DB2.LE.DB1) B1=BET
IF (DB3.GE.LE2.AND.DB2.LE.DB1) LEN=1
IF (DB3.GE.LE2.AND.DB2.LE.DB1) FBT=FBT1+2.*(FBT2+FBT3)
GO TO 26
B1=2.0*(BET1+BT2)
LEN=1.0*(FBT1+FBT2+FBT3)
FBT=2.0*(FBT1+FBT2+FBT3)
WRITE (6,401) J,B1,FBT,FBT1,FBT2,FBT3,TX1,YN1
IF (CKZERO (H2-H1)) GO TO 26
IP=103
IF (NP1.EQ.1) J1=J1+1
SPHI=SIN (ANGLE)
IF (Z (J1+1).LE.H2) CALL POINT (H2,YN,N,NP,TX,IP)
IF (Z (J1+1).LE.H2) J2=J1

```

19

20

*

23

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RN=TX 1/YN1
IF (SPHI.GE.EN) RN=1.
SPHI=SPHI/RN
THET=ARSIN (SPHI)
GO TO 5
THET=ANGLE+ (ETA-B 1) / (1.+FBT/TANG)
DB1=B 1/CA
E=BET 1/CA
TH1=THET/CA
WRITE(6,404) E1,DB1,FBT,TH1,TANG
IF (THET.GT.IN.OR.THET.LT.TM) THET=(TN+TM) / 2.
TH1=THET/CA
WRITE(6,404) BET1,B,FBT,TH1
TN1=TN/CA
TH1=TM/CA
WRITE(6,405) IN,TM,TN1,TM1
SPHI= SIN (THET)
TANG= TAN (THET)
IF (ABS (BETA-B1) .LT. 1.E-7.OR.ABS (ANGLE-THET) .LT. 1.E-7) GO TO 28
CONTINUE
THET=(ANGLE+THET) / 2.
ANGLE=THET/CA
IF (BETA.LE.0.0) H1=H2
WRITE(6,406) ANGLE,ITER
RETURN

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LL 13010
LL 13020
LL 13030
LL 13040
LL 13050
LL 13060
LL 13070
LL 13080
LL 13090
LL 13100
LL 13110
LL 13120
LL 13130
LL 13140
LL 13150
LL 13160
LL 13170
LL 13180
LL 13190
LL 13200
LL 13210
LL 13220
LL 13230
LL 13240
LL 13250
LL 13260
LL 13270
THLL 13280
LL 13290
LL 13300
SL 13310
LL 13320
LL 13330

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101

28

C 401

404

405

406

C

C

C

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SUBROUTINE ARRAY(P,T,WH,WO,VX,C1,C2,C7,C7A)
COMMON /LOWTRN/ IATM,NL,HZ1(34),HZ2(34),Z0(34),P01(34),P02(34),
*P03(34),P04(34),P05(34),P06(34),WH01(34),WH02(34),WH03(34),WH04(34),T01(34),T02(34),T03(34),T04(34),
*T05(34),T06(34),WO01(34),WO02(34),WO03(34),WO04(34),WO05(34),WO06(34),
*WHC6(34),FO(67),FO01(190),FO02(190),FO03(190),FO04(190),
*TR(67),FO05(190),FO06(190),FO07(190),FO08(190),FO09(190),FO10(190),
*C105(190),C106(190),C107(190),C108(190),C109(190),C110(190),
*C111(190),C112(190),C113(190),C114(110),C201(190),C202(190),
*C203(190),C204(190),C205(190),C3(540),C4(133),C5(15),C701(45),
*C702(45),C703(45),C704(45),C7A01(45),C7A02(45),C7A03(45),

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LL 13340
LL 13350
LL 13360
LL 13370
LL 13380
LL 13390
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LL 13970
LL 13980
LL 13990

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* C7A04 (45) , C8 (102) , VX01 (45) , VX02 (45) , VX03 (45) , VX04 (45) ,
* EH (10, 34) , M, H1, M2, M3, RE, CM, CO, PI, CA, REARTH (7) , Z (34) ,
* DIMENSION P (7, 34) , T (7, 34) , WH (7, 34) , WO (7, 34) , VX (4, 45) ,
* C1 (2580) , C2 (1575) , C7 (4, 45) , C7A (4, 45)

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CC

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DO 10 I=1, 34
P (1, I) = P02
P (2, I) = P03
P (3, I) = P04
P (4, I) = P05
P (5, I) = P06
P (6, I) = T01
T (1, I) = T02
T (2, I) = T03
T (3, I) = T04
T (4, I) = T05
T (5, I) = T06
T (6, I) = WH01
WH (1, I) = WH02
WH (2, I) = WH03
WH (3, I) = WH04
WH (4, I) = WH05
WH (5, I) = WH06
WH (6, I) = WO01
WO (1, I) = WO02
WO (2, I) = WO03
WO (3, I) = WO04
WO (4, I) = WO05
WO (5, I) = WO06
CONTINUE
DC 20 I=1, 45
VX (1, I) = VX01
VX (2, I) = VX02
VX (3, I) = VX03
VX (4, I) = VX04
CONTINUE
DO 20 I=1, 45
C7 (1, I) = C701
C7 (2, I) = C702
C7 (3, I) = C703
C7 (4, I) = C704
C7A (1, I) = C7A01
C7A (2, I) = C7A02
C7A (3, I) = C7A03
C7A (4, I) = C7A04

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10

20

LL 01280
LL 01390
LL 13480

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LL 13750
LL 13760
LL 13780

LL 13640
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LL 13660
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LL 13640
LL 13650
LL 13660
LL 13670


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C C C      DATA      PI /3.14159265358979/
*** THE CCNVERSION FROM DEGREES TO RADIAN.
C C C      DATA      CA /0.017453292519943/
*** THE MEAN RADIUS OF THE EARTH (INTERNATIONAL SPHERIOD) IN KM.
C C C C C    DATA      REBAR /6371.299315/
*** THE POLAR RADIUS OF THE EARTH (INTERNATIONAL SPHERIOD) IN KM.
C C C C C    DATA      REPOL /6356.911946/
*** THE EQUATORIAL RADIUS OF THE EARTH (INTERNATIONAL SPHERIOD) IN KM.
C C C C C    DATA      REEQU /6378.388000/
*** VALUES OF THE RADIUS OF THE EARTH FOR THE MODEL ATMOSPHERES IN KM.
C C C C C    DATA      REARTH {1} /6376.951637/
C C C C C    DATA      REARTH {2} /6367.659029/
C C C C C    DATA      REARTH {3} /6367.659029/
C C C C C    DATA      REARTH {4} /6362.287754/
C C C C C    DATA      REARTH {5} /6362.287754/
C C C C C    DATA      REARTH {6} /6367.659029/
C C C C C    DATA      REARTH {7} /6371.299315/
*** NUMBER OF TYPES OF MODEL ATMOSPHERES
C C C      DATA      IATH /6/
*** NUMBER OF LEVELS IN MODEL ATMOSPHERE DATA
C C C      DATA      NL /34/
*** HAZE MODEL' 23 KM. VISIBLE RANGE
C C C      DATA      HZ1 /2.830E+03,1.245E+03,5.374E+02,2.257E+02,
* 1.193E+02,8.992E+01,6.341E+01,5.893E+01,6.073E+01,5.822E+01,
* 5.679E+01,5.320E+01,5.589E+01,5.159E+01,5.052E+01,4.747E+01,
* 4.514E+01,4.460E+01,4.317E+01,3.636E+01,2.669E+01,1.935E+01,
* 1.456E+01,1.114E+01,8.831E+00,7.434E+00,2.239E+00,5.893E-01,
* 1.551E-01,4.084E-02,1.078E-02,5.553E-05,1.970E-08,0.000E+00/
*** HAZE MODEL' 5 KM. VISIBLE RANGE
C C C

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LL 14190
LL 14200
LL 14210
LL 14220
LL 14230
LL 14240
LL 14250
LL 14260
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LL 14290
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LL 14320
LL 14330
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LL 14570
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LL 14590
LL 14600
LL 14610
LL 14620
LL 14630
LL 14640
LL 14650

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DATA 4.54E+02,8.992E+01,6.341E+01,5.893E+03,1.845E+03,6.735E+02,
* 2.679E+01,5.320E+01,5.589E+01,5.159E+01,5.052E+01,4.747E+01,
* 4.514E+01,4.460E+01,4.317E+01,3.636E+01,2.669E+01,1.935E+01,
* 1.456E+01,1.114E+01,8.331E+00,7.434E+00,2.239E+00,5.893E-01,
* 1.551E-01,4.084E-02,1.078E-02,5.553E-05,1.970E-08,0.000E+00/
C *** ALTITUDE (KM.) AT LEVEL I
C
DATA 9.0,10.0,11.0,12.0,13.0,14.0,15.0,16.0,17.0,18.0,19.0,20.0,
* 21.0,22.0,23.0,24.0,25.0,30.0,35.0,40.0,45.0,50.0,70.0,
* 100.0,99999.0/
C *** PRESSURE (MB.) AT LEVEL I FOR THE MODEL ATMOSPHERES
C
DATA 7.150E+02,6.330E+02,5.590E+02,4.920E+02,4.320E+02,3.780E+02,
* 3.290E+02,2.860E+02,2.470E+02,2.130E+02,1.820E+02,1.560E+02,
* 1.320E+02,1.110E+02,9.370E+01,7.890E+01,6.660E+01,5.650E+01,
* 4.800E+01,4.090E+01,3.500E+01,3.000E+01,2.570E+01,1.220E+01,
* 6.000E+00,3.050E+00,1.590E+00,8.540E-01,5.790E-02,3.000E-04,
* 0.000E+00/
DATA 7.100E+02,6.280E+02,5.540E+02,4.870E+02,4.260E+02,3.720E+02,
* 3.240E+02,2.810E+02,2.430E+02,2.090E+02,1.790E+02,1.530E+02,
* 1.300E+02,1.110E+02,9.500E+01,8.120E+01,6.950E+01,5.950E+01,
* 5.100E+01,4.370E+01,3.760E+01,3.220E+01,2.770E+01,1.320E+01,
* 6.520E+00,3.330E+00,1.760E+00,9.510E-01,6.710E-02,3.000E-04,
* 0.000E+00/
DATA 6.938E+02,6.081E+02,5.313E+02,4.627E+02,4.018E+02,3.497E+02,
* 2.992E+02,2.568E+02,2.199E+02,1.882E+02,1.610E+02,1.378E+02,
* 1.178E+02,1.007E+02,8.610E+01,7.350E+01,6.280E+01,5.370E+01,
* 4.580E+01,3.910E+01,3.340E+01,2.860E+01,2.430E+01,1.110E+01,
* 5.180E+00,2.530E+00,1.290E+00,6.820E-01,4.670E-02,3.000E-04,
* 0.000E+00/
DATA 7.000E+02,6.160E+02,5.410E+02,4.730E+02,4.100E+02,3.590E+02,
* 3.107E+02,2.677E+02,2.300E+02,1.977E+02,1.700E+02,1.460E+02,
* 1.250E+02,1.080E+02,9.280E+01,7.980E+01,6.860E+01,5.890E+01,
* 5.070E+01,4.360E+01,3.750E+01,3.227E+01,2.780E+01,1.340E+01,
* 6.610E+00,3.400E+00,1.810E+00,9.870E-01,7.070E-02,3.000E-04,
* 0.000E+00/
DATA 6.798E+02,5.932E+02,5.158E+02,4.467E+02,3.853E+02,3.308E+02,
* 2.829E+02,2.418E+02,2.067E+02,1.766E+02,1.510E+02,1.291E+02,
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LL 14660
LL 14670
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LL 15100
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LL 15120

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* * *	1	.103E+02,	9	.431E+01,	8	.050E+01,	6	.882E+01,	5	.875E+01,	5	.014E+01,	
* * *	4	.277E+01,	3	.647E+01,	3	.109E+01,	2	.649E+01,	2	.256E+01,	1	.020E+01,	
* * *	4	.701E+00,	2	.243E+00,	1	.113E+00,	5	.719E-01,	4	.016E-02,	3	.000E-04,	
* * *	0	.000E+00,	/										
	DATA		P06										
* * *	7	.012E+02,	6	.166E+02,	5	.405E+02,	/	1	.013E+03,	8	.986E+02,	7	.950E+02,
* * *	3	.080E+02,	2	.650E+02,	2	.270E+02,	1	.940E+02,	1	.658E+02,	1	.417E+02,	
* * *	1	.211E+02,	1	.035E+02,	8	.850E+01,	7	.565E+01,	6	.467E+01,	5	.529E+01,	
* * *	4	.729E+01,	4	.047E+01,	3	.467E+01,	2	.972E+01,	2	.549E+01,	1	.197E+01,	
* * *	5	.746E+00,	2	.871E+00,	1	.491E+00,	7	.978E-01,	5	.520E-02,	3	.008E-04,	
* * *	0	.000E+00,	/										

DATA P07 /34*0.0/

*** TEMPERATURE (K) AT LEVEL I FOR THE MODEL ATMOSPHERES

DATA	T01	270.0	264.0	257.0	250.0	244.0	/300.0	294.0	0.0	2245.0	288.0	284.0	277.0
*		204.0	197.0	195.0	199.0	203.0	237.0	230.0	0.0	2215.0	217.0	217.0	210.0
*		221.0	232.0	243.0	254.0	265.0	270.0	211.0	0.0	2110.0	217.0	217.0	219.0
DATA	T02	267.0	261.0	255.0	248.0	242.0	/294.0	290.0	0.0	2285.0	2279.0	2279.0	273.0
*		216.0	216.0	216.0	216.0	217.0	235.0	229.0	0.0	2222.0	2216.0	2216.0	216.0
*		224.0	234.0	245.0	258.0	270.0	276.0	218.0	0.0	2210.0	2220.0	2220.0	223.0
DATA	T03	249.7	243.7	237.7	231.7	225.7	/272.2	268.7	2.2	2165.7	2261.7	2261.7	255.7
*		217.2	216.7	216.2	215.7	215.2	215.7	219.2	2.2	2115.2	2218.2	2218.2	217.7
*		215.2	217.4	227.8	243.2	258.5	267.5	230.7	2.2	2110.7	2215.0	2215.0	215.2
DATA	T04	260.0	253.0	246.0	239.0	232.0	/287.0	282.0	0.0	2276.0	2271.0	2271.0	266.0
*		225.0	225.0	225.0	225.0	225.0	225.0	225.0	0.0	2225.0	2225.0	2225.0	225.0
*		228.0	235.0	247.0	262.0	274.0	277.0	216.0	0.0	2210.0	2210.0	2210.0	226.0
DATA	T05	240.9	234.1	227.3	220.6	217.2	/257.1	259.1	1.2	2155.1	2252.7	2252.7	247.7
*		217.2	216.6	216.0	215.4	214.8	217.2	217.2	1.2	2113.0	2217.2	2217.2	217.2
*		211.2	216.0	222.2	234.7	247.0	259.3	245.3	1.2	2110.7	2212.4	2212.4	211.8
DATA	T06	255.7	249.2	242.7	236.2	229.7	/288.1	281.1	2.2	2175.6	2268.7	2268.7	262.2
*		216.6	216.6	216.6	216.6	216.6	223.6	216.6	2.2	2118.6	2216.6	2216.6	216.6
*		221.6	226.5	236.5	253.4	264.2	270.2	219.6	2.2	2110.6	2210.6	2210.6	220.6

DATA T07 /34*0.0/

LL	15130
LL	15140
LL	15150
LL	15160
LL	15170
LL	15180
LL	15190
LL	15200
LL	15210
LL	15220
LL	15230
LL	15240
LL	15250
LL	15260
LL	15270
LL	15280
LL	15290
LL	15300
LL	15310
LL	15320
LL	15330
LL	15340
LL	15350
LL	15360
LL	15370
LL	15380
LL	15390
LL	15400
LL	15410
LL	15420
LL	15430
LL	15440
LL	15450
LL	15460
LL	15470
LL	15480
LL	15490
LL	15500
LL	15510
LL	15520
LL	15530
LL	15540
LL	15550
LL	15560
LL	15570
LL	15580
LL	15590

**	-1.10,	-1.13,	-1.19,	-1.22,	-1.28,	-1.30,	-1.33,	-1.36,	-1.39,	-1.43,	184230
	-1.48,	-1.50,	-1.52,	-1.57,	-1.61,	-1.66,	-1.70,	-1.72,	-1.78,	-1.81,	18430
	DATA	C107									18440
**	-1.89,	-1.92,	0.03,	0.08,	0.10,	0.16,	0.20,	0.24,	0.28,	0.31,	184450
	-2.61,	-2.71,	-2.83,	-2.95,	-3.00,	-3.05,	-3.10,	-3.15,	-3.20,	-3.25,	18450
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18460
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18470
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18480
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18490
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18500
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18510
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18520
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18530
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18540
*	-3.78,	-3.33,	-3.01,	-2.82,	-2.68,	-2.49,	-2.30,	-2.13,	-2.00,	-1.81,	18550
*	-1.60,	-1.41,	-1.13,	-0.90,	-0.79,	-0.63,	-0.49,	-0.36,	-0.28,	-0.16,	18560
*	-0.06,	0.08,	0.20,	0.28,	0.41,	0.54,	0.69,	0.80,	0.92,	1.04,	18570
*	1.19,	1.19,	1.42,	1.43,	1.70,	1.62,	1.54,	1.41,	1.23,	1.08,	18580
*	1.96,	1.97,	2.02,	2.01,	1.94,	1.84,	1.73,	1.61,	1.53,	1.38,	18590
*	2.30,	2.16,	2.02,	2.02,	2.16,	2.13,	2.03,	1.90,	1.71,	1.56,	18600
*	2.56,	2.51,	2.39,	2.49,	2.62,	2.57,	2.41,	2.28,	2.10,	1.93,	18610
*	2.45,	2.51,	2.23,	2.49,	2.52,	2.51,	2.37,	2.21,	2.03,	1.86,	18620
	DATA	C108									18630
**	2.51,	2.70,	2.76,	2.84,	2.92,	3.00,	3.07,	3.14,	3.20,	3.26,	18640
	-2.81,	-2.72,	-2.53,	-2.38,	-2.25,	-2.11,	-1.98,	-1.84,	-1.70,	-1.56,	18650
*	3.26,	3.06,	3.34,	3.28,	3.37,	3.44,	3.51,	3.57,	3.63,	3.69,	18660
*	3.07,	3.07,	3.31,	3.21,	3.31,	3.37,	3.42,	3.47,	3.51,	3.55,	18670
*	3.39,	3.39,	3.33,	3.29,	3.33,	3.38,	3.42,	3.46,	3.50,	3.53,	18680
*	3.40,	3.38,	3.39,	3.27,	3.26,	3.24,	3.22,	3.20,	3.18,	3.16,	18690
*	3.66,	3.75,	3.78,	3.56,	3.52,	3.48,	3.44,	3.40,	3.36,	3.32,	18700
*	3.91,	3.62,	3.64,	3.53,	3.52,	3.51,	3.50,	3.49,	3.48,	3.47,	18710
*	1.68,	1.62,	1.64,	1.56,	1.52,	1.51,	1.50,	1.49,	1.48,	1.47,	18720
*	1.40,	1.41,	1.43,	1.40,	1.42,	1.44,	1.45,	1.46,	1.47,	1.48,	18730
*	1.09,	1.16,	1.21,	1.20,	1.22,	1.24,	1.25,	1.26,	1.27,	1.28,	18740
*	1.10,	1.10,	1.09,	1.10,	1.11,	1.11,	1.11,	1.10,	1.09,	1.08,	18750
*	0.90,	0.90,	0.86,	0.87,	0.88,	0.88,	0.88,	0.87,	0.86,	0.85,	18760
*	0.42,	0.31,	0.20,	0.11,	0.08,	0.07,	0.06,	0.05,	0.04,	0.03,	18770
*	-0.63,	-0.73,	-0.84,	-0.93,	-1.04,	-1.14,	-1.24,	-1.34,	-1.44,	-1.54,	18780
*	-1.64,	-1.74,	-1.84,	-1.94,	-2.04,	-2.14,	-2.24,	-2.34,	-2.44,	-2.54,	18790
*	-2.64,	-2.74,	-2.84,	-2.94,	-3.04,	-3.14,	-3.24,	-3.34,	-3.44,	-3.54,	18800
	DATA	C109									18810
*	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18820
	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18830
	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18840
	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18850
	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18860
	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18870
	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	-5.00,	18880

[illegible]

DATA

C110

[illegible]

DATA

C111

DATA	C111	1.58	1.62	1.61	1.62	1.66	1.67	1.63	1.71	1.72	1.70
1.58	1.67	1.62	1.61	1.61	1.62	1.66	1.67	1.63	1.71	1.72	1.70
1.70	1.67	1.63	1.66	1.70	1.67	1.66	1.67	1.56	1.49	1.42	1.38
1.26	1.20	1.13	1.14	1.19	1.29	1.16	1.29	1.50	1.72	1.86	1.78
1.82	1.89	1.82	1.89	1.99	2.09	1.81	2.09	2.19	2.50	2.36	2.20
1.98	1.90	1.83	1.88	1.72	1.69	1.88	1.69	0.52	0.53	0.40	0.41
0.98	0.63	0.43	0.29	0.16	0.05	0.29	0.05	0.02	0.03	0.00	0.01

82

CCNTINUED DATA PACKAGE

[illegible]

0.009338	0.008227	0.01010	0.00923	0.01010	0.00878
0.008223	0.00808	0.00000	0.22026	0.17989	0.15800
0.038223	0.32979	0.28540	0.03126	0.02510	0.02068
0.09151	0.07078	0.04184	0.01533	0.01479	0.01389
0.01767	0.01699	0.01654	0.01199	0.02112	0.01870
0.01102	0.01019	0.01778	0.01514	0.01455	0.01339
0.01744	0.01714	0.01588	0.01384	0.01480	0.01381
0.01286	0.01368	0.01384	0.01384	0.01353	0.01427
0.01302	0.01204	0.00000	0.21638	0.17869	0.15800
0.35815	0.31452	0.27624	0.03574	0.02975	0.02451
0.09459	0.07434	0.04594	0.01693	0.01618	0.01518
0.02027	0.01940	0.01876	0.01947	0.02041	0.02121
0.01217	0.01141	0.01771	0.01509	0.01454	0.01367
0.01719	0.01690	0.01577	0.01431	0.01317	0.01341
0.01289	0.01350	0.01355	0.01431	0.01317	0.01323
0.01249	0.01153	0.00000	0.16615	0.16065	0.15800
0.19374	0.18391	0.17637	0.12826	0.12323	0.1186
0.14855	0.14897	0.13491	0.07573	0.06413	0.05932
0.10644	0.09893	0.09340	0.04619	0.04556	0.04282
0.04383	0.04239	0.04346	0.02729	0.02819	0.03378
0.03564	0.03344	0.02846	0.04481	0.04295	0.03973
0.04014	0.04205	0.04368	0.01450	0.01050	0.01000
0.03166	0.02641	0.00000	0.00924	0.00348	0.00369
0.09530	0.05660	0.02060	0.00171	0.00154	0.00248
0.00914	0.01020	0.01080	0.00629	0.00504	0.00702
0.00487	0.00360	0.00222	0.00937	0.00698	0.00549
0.00295	0.00360	0.00423	0.00607	0.00506	0.00587
0.01180	0.01310	0.01430	0.01317	0.01114	0.01095
0.00386	0.00464	0.00691	0.00767	0.00677	0.00767
0.00562	0.00581	0.00000	0.01317	0.01114	0.01095
0.07945	0.03661	0.02110	0.00700	0.00437	0.00463
0.01058	0.01070	0.00933	0.00321	0.00388	0.00462
0.00250	0.00214	0.00232	0.01126	0.01209	0.01378
0.00617	0.00807	0.01254	0.00570	0.00535	0.00516
0.00832	0.00810	0.00680	0.00767	0.00677	0.00767
0.00538	0.00834	0.00696	0.00767	0.00677	0.00767
0.00749	0.00761	0.00000	0.03011	0.02640	0.02496
0.08865	0.05331	0.03938	0.01272	0.00998	0.00920
0.02050	0.01905	0.01548	0.00575	0.00608	0.00649
0.00622	0.00569	0.00563	0.01141	0.01207	0.01344
0.00729	0.00884	0.01253	0.00646	0.00612	0.00587
0.00879	0.00858	0.00743	0.00770	0.00689	0.00759
0.00595	0.00837	0.00714	0.00770	0.00689	0.00759
0.00729	0.00730	0.00000	0.00383	0.00223	0.00183
0.01647	0.00710	0.00383	0.00161	0.00243	0.00369
0.00169	0.00191	0.00200	0.00629	0.00554	0.00521
0.00711	0.00298	0.00344	0.00629	0.00554	0.00521

90

09	18	63	18	00	00	00	00	00	00	22	28	90	00	40	39	78	12	24	20	20
32	24	10	15	55	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
92	00	87	20	00	00	00	00	00	00	47	17	07	10	00	35	12	20	21	83	0
10	00	93	30	00	00	00	00	00	00	36	55	08	40	00	33	16	30	86	97	0
33	25	10	15	55	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
93	60	20	65	10	00	00	00	00	00	10	90	00	40	00	08	14	15	00	23	30
81	00	20	65	10	00	00	00	00	00	54	05	00	40	00	00	14	15	00	50	10
23	25	20	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
11	90	18	30	00	00	00	00	00	00	88	08	60	10	00	38	38	00	69	17	80
93	20	51	83	10	00	00	00	00	00	61	08	10	50	00	40	39	00	13	14	0
23	25	20	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
93	10	47	19	80	00	00	00	00	00	97	00	39	00	19	00	40	18	00	32	10
95	40	78	15	80	00	00	00	00	00	90	00	64	05	50	79	18	00	32	27	0
23	25	20	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
93	20	28	18	30	00	00	00	00	00	10	00	32	00	13	00	33	00	17	34	00
07	30	00	98	13	00	00	00	00	00	31	00	10	00	52	00	70	60	50	17	00
33	25	30	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
92	19	30	11	60	00	00	00	00	00	17	80	62	00	58	00	30	14	00	82	00
69	21	89	16	20	00	00	00	00	00	10	90	00	40	00	08	14	15	00	23	30
33	24	31	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
11	89	27	21	00	00	00	00	00	00	20	17	48	00	60	00	78	80	20	14	02
43	23	31	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
22	36	99	38	10	00	00	00	00	00	50	88	00	70	00	26	00	80	33	96	37
56	39	89	30	10	00	00	00	00	00	20	43	00	94	00	10	60	10	53	96	15
43	22	31	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5
81	36	57	18	10	00	00	00	00	00	23	68	42	00	21	00	51	29	11	24	85
43	22	31	00	45	55	55	55	55	55	34	42	54	35	41	12	44	00	00	00	5

APPENDIX C
SAMPLE OUTPUT

SUB-ARCTIC (60 DEG. LAT.) WINTER MODEL ATMOSPHERIC
CONTINENTAL AEFCSOL MODEL

HAZE MODEL = 23.0 KM VISUAL RANGE AT SEA LEVEL

H1 = 2.500 KM, H2 = 8.500 KM, ANGLE = 65.0000 GEOM. RANGE = 14.15 KM, BETA = 0.11537 DEG

SLANT PATH BETWEEN ALTITUDES H1 AND H2 WHERE H1 = 2.500 KM H2 = 8.500 KM, ZENITH ANGLE = 65.000 DEGREES
FREQUENCY RANGE V1 = 2350.0 CM-1 TO V2 = 2450.0 CM-1 FOR DV = 5.0 CM-1 (4.08 - 4.26 MICRONS)

AD-A132 123

CALCULATION OF ATMOSPHERIC TRANSMITTANCE BY IBM 3033
COMPUTER CODE LOWTRAN IIIB(U) NAVAL POSTGRADUATE SCHOOL
MONTEREY CA M SHIN JUN 83

2/2

UNCLASSIFIED

F/G 20/6

NL





MICROCOPY RESOLUTION TEST CHART
NATIONAL BUREAU OF STANDARDS-1963-A

APPENDIX D DEFINITIONS AND SYMBOLS

AB	Absorption at frequency ν ; also average transmittance
AH21, AH22	Aerosol number density
AJ	Equivalent absorber amount per km at level J
ALP	Angle of arrival at adjacent level
ANGLE	Input zenith angle (degree)
BET	Angle subtended at the earth's center as path traverses adjacent levels
BETA	Total angle subtended path at earth's center
BJ	Equivalent absorber amount per km at level J+1
CA	Conversion factor from degree to radians
CO	Wavelength dependent coefficient used in refractive index expression
CW	Wavelength dependent coefficient used in refractive index expression
C1	Log absorption coefficient for water vapor
C2	Log absorption coefficient for uniformly mixed gas
C3	Log absorption coefficient for ozone
C4	Absorption coefficient for nitrogen
C5	Absorption coefficient for water vapor continuum
C6	Extinction coefficient for molecular scattering
C7	Extinction coefficient for aerosol models
C7A	Aerosol absorption coefficient
C8	Absorption coefficient for ozone (UV and visible region)
D	water vapor amount (pr.cm/km) at level I
DP	Dew point temperature (C)
DS	Path length from level I to level I + 1
DV	Wave number increment at which transmittance is calculated
DZ	Height increment from level I to I + 1
E(K)	Equivalent absorber amount per km at height H1
EH(1,I)	Equivalent absorber amount per km for water vapor at level Z(I)
EH(2,I)	Equivalent absorber amount per km for carbon dioxide, etc. at level Z(I)
EH(3,I)	Equivalent absorber amount per km for ozone at level Z(I)
EH(4,I)	Equivalent absorber amount per km for nitrogen at level Z(I)
EH(5,I)	Equivalent absorber amount per km for water vapor continuum at level Z(I)
EH(6,I)	Equivalent absorber amount per km for molecular scattering at level Z(I)
EH(7,I)	Equivalent absorber amount per km for waerosol extinction at level Z(I)
EH(8,I)	Equivalent absorber amount per km for ozone (UV and visible) at level Z(I)
EH(9,I)	Mean refractive index of layer above level Z(I)
EV	Integrated absorber amount from level I to I + 1
FAC	Factor for exponential and linear interpolation
FO	Transmissiön function logarithmic absorber amount scale for ozone
FW	Transmissiön function logarithmic absorber amount scale for water vapor and uniformly mixed gases
H	Altitude (km)
H1	Initial Altitude (km)
H2	Final Altitude (km)

HAZE	Aerosol number density (no. cm-3)
HM	Estimated tangent height (km)
HMIN	Minimum altitude of path trajectory (km)
HZ1	Aerosol number density (no. cm-3) for 23 km visual range
HZ1	Aerosol number density (no. cm-3) for 5 km visual range
I	Running integer used as altitude indicator
IAERO	Indicator for type of aerosol model
IATOM	Number of levels in model atmosphere
IDV	Frequency increment (cm-1)
IFIND	Indicator for using subroutine ANGL
IHAZE	Aerosol model indicator
IM	Parameter used when reading in a new atmospheric model
IP	Indicator for using subroutine POINT to calculate reference index only (IP = 0)
IR	Card printer number
ITER, ITES	Iteration counters
ITYPE	Indicator for type of atmospheric path
IV	Frequency at which transmittance is calculated
IV1	Starting frequency
IV2	Last frequency
IW	Line printer number
IXY	Parameter for terminating program and cycling indicator
J	Running integer for altitude identification
JMIN	Altitude indicator for minimum height of path
JP	Print option for altitude H1
J1	Level indicator for altitude H1
J2	Level indicator for altitude H2
K	Absorber indicator, K=1, 2, 3, etc., corresponds to water vapor, uniformly mixed gases, ozone, etc., respectively
K2	Cycling parameter for downward looking paths
L	Frequency indicator for ozone transmittance calculation
LEN	Parameter used for defining longest of two paths
M	Integer used to identify required model atmosphere
ML	Number of levels in radiosonde data input (MODEL7)
MODEL	Integer used to identify required model atmosphere
M1	Integer for selecting temperature altitude profile for (M=M1)
M2	Integer for selecting water vapor altitude profile for (M=M2)
M3	Integer for selecting ozone altitude profile for (M=M3)
N	Indicator for level below given input altitude used in PCINT subroutine
NL, NLF	Number of levels in model atmosphere data
NP	Indicator for determining whether H1 or H2 coincide with levels in the model atmosphere
NP1	Value of NP for altitude H1
NP2	Value of NP for altitude H2
NS1, NS2	Counters corresponding to WS1, WS2
P(M, 1)	Pressure (mb) at level I for model atmosphere M
PHI	Angle of arrival at H2
PPW	Partial pressure of water vapor (in atmosphere)
PS	Total pressure in atmospheres
PSI	Angular deviation of path from initial direction
PT	Product of total pressure (atm) and the square root of $273/T(M, I)$
RANGE	Path length (km)
RE, REATH	Earth radius (km)
REF	Refractive index of air at level I
RH	Relative humidity (%)
RN	Ratio of refractive indices of air above and

below a given level

RX Ratio of earth center distances between adjacent levels

RO Earth radius (km) read in as input (=RE)

R1 The product of the sine of the initial zenith angle and the earth center distance to starting altitude

SALP Sine of angle of arrival at adjacent level

SPHI Sine of the local zenith angle at a given level

SR Slant range (km)

SUM Accumulated absorption

T(M, I) Temperature (K) for model atmosphere M at level I

THET Zenith angle at a given level (in radians)

THETA Zenith angle at a given level (in degrees)

TMP Ambient temperature (C)

TR Transmittance scales for transmission functions

TS Ratio of standard temperature (273.15 K) to temperature level I

TS1 Ratio of 296 K to temperature at level I

TT Ratio 273.15/(TMP + 273.15)

TX(K) Equivalent absorber amounts per km at a given altitude obtained from POINT; also transmittance values at a given wavelength for each absorber type (K = 1,8)

TX(9) Total transmittance at frequency IV

TX(10) Absorption due to aerosol only at frequency IV

TX1 Refractive index of layer above initial alt. H1

TX2 Refractive index of layer above final alt. H2

TX1 Refractive index of layer above minimum alt. HMIN

VH(K) Integral of the equivalent absorber amounts from H1 to level I

VIS Visual range (km) at sea level

VX Wavelength at which aerosol coefficients are read in (micrometer)

V1 Initial frequency for transmittance calculation

V2 Final frequency for transmittance calculation

W(K) Total equivalent absorber amount for entire path

WH(M, I) Water vapor density for atmospheric model M at level I (gm m⁻³)

WL, WL1, WL2 Wavelength in microns

WO(M, I) Ozone density for atmospheric model M at level I

WS1 Transmission function scaling factor for water vapor at given wavelength

WS2 Transmission function scaling factor for carbon dioxide, etc., at given wavelength

WS3 Transmission function scaling factor for ozone at given wavelength

X Input height to POINT subroutine

XI Wavenumber interpolation parameter

XX Wavenumber identification parameter for UV ozone transmittance calculation

X1 Earth center distance of level I

X2 Earth center distance of level I + 1

YN Refractive index of layer below input height from POINT subroutine

YN1 Refractive index of layer below initial alt. H1

YN2 Refractive index of layer below final alt. H2

YY Aerosol absorption coefficient at frequency V

Z(I), ZO(I) Altitude at level I in km

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